

Predictability of Atlantic Overturning Circulation and Associated Surface Patterns in Two CCSM3 Large Ensemble Experiments

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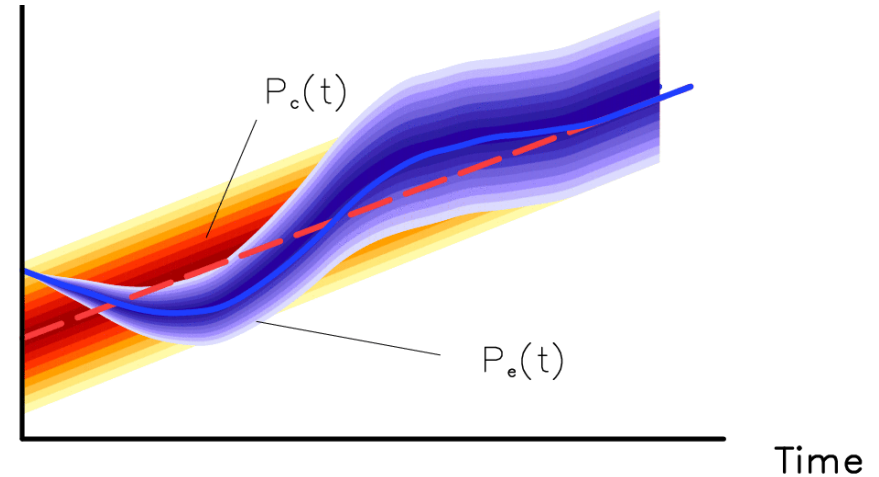
June 29, 2011

Predictability

Total predictability: $P_e(t)$ vs. $P_c(0)$

Initial-value predictability: $P_e(t)$ vs. $P_c(t)$

Forced predictability: $P_c(t)$ vs. $P_c(0)$



Kleeman (2002)

$$R = \int_S P_x(s) \log_2 \left[\frac{P_x(s)}{P_b(s)} \right] ds$$

$P_x(s)$: prediction
 $P_b(s)$: base

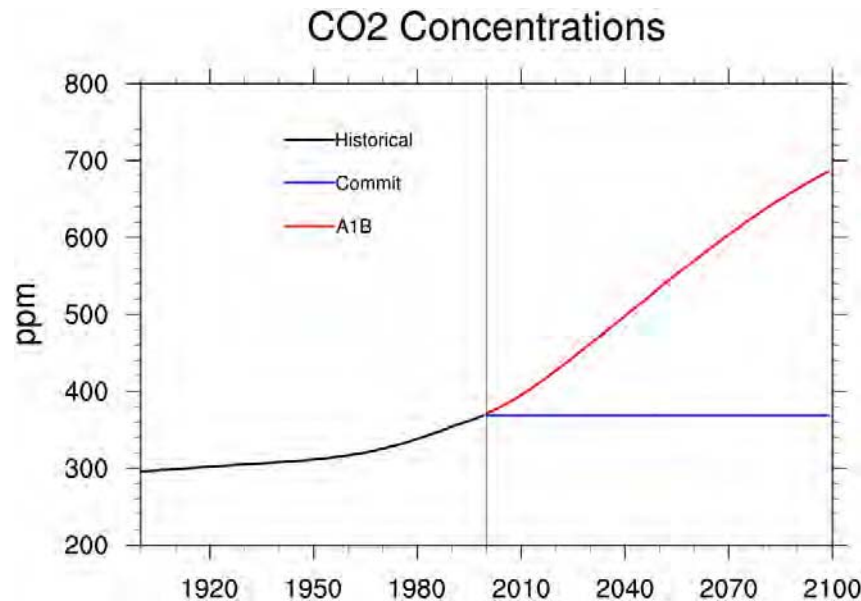
For normal distribution:

$$R = \frac{1}{2} \log_2(e) \left\{ \underbrace{\ln \left[\frac{\det(\sigma_b^2)}{\det(\sigma_x^2)} \right]}_{\text{dispersion}} + \underbrace{\text{trace} \left(\frac{\sigma_x^2}{\sigma_b^2} \right)}_{\text{signal}} + \underbrace{(\mu_x - \mu_b)^T (\sigma_b^2)^{-1} (\mu_x - \mu_b) - n}_{\text{bias}} \right\}$$

dispersion

signal

Model & Experiments



- CCSM3.0, T42 atm, 1deg ocn
- 40-member ensemble run
 - A1B, Commitment
 - 2000-2061
 - atm perturbed ICs
 - Same ocn/Ind/ice ICs
- A1B (II), A1B(III), A1B (IV)
 - Each has unique ocn/ice/Ind ICs
 - 20 years
- Last 700 yrs of the 1000-yr control

Initial-value Predictability & Forced Predictability

Assumptions:

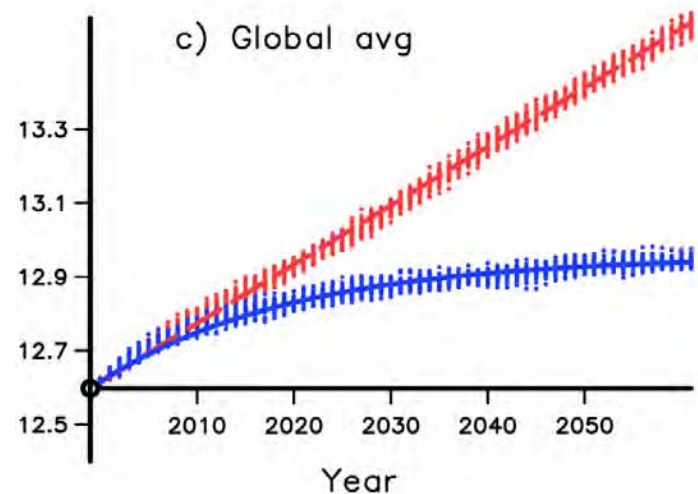
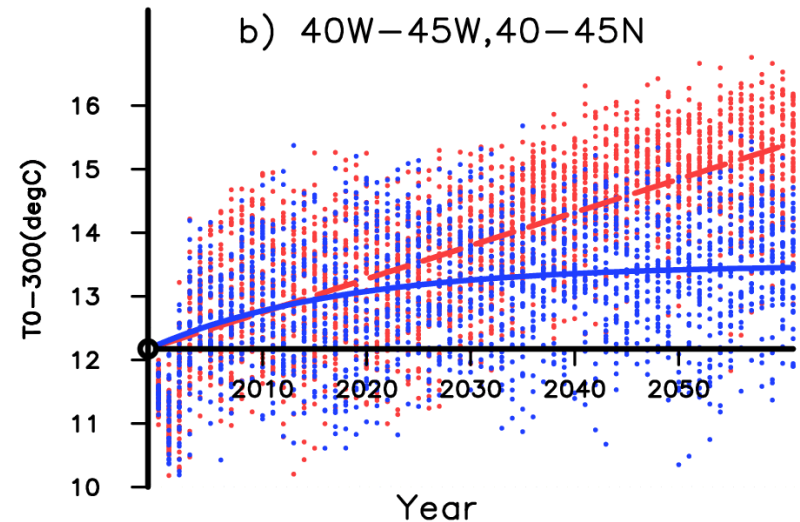
1. Gaussians
2. Climate covariance does not change with time.
3. Climate mean can be well approximated by an analytical function of time.

Time evolving climate mean:

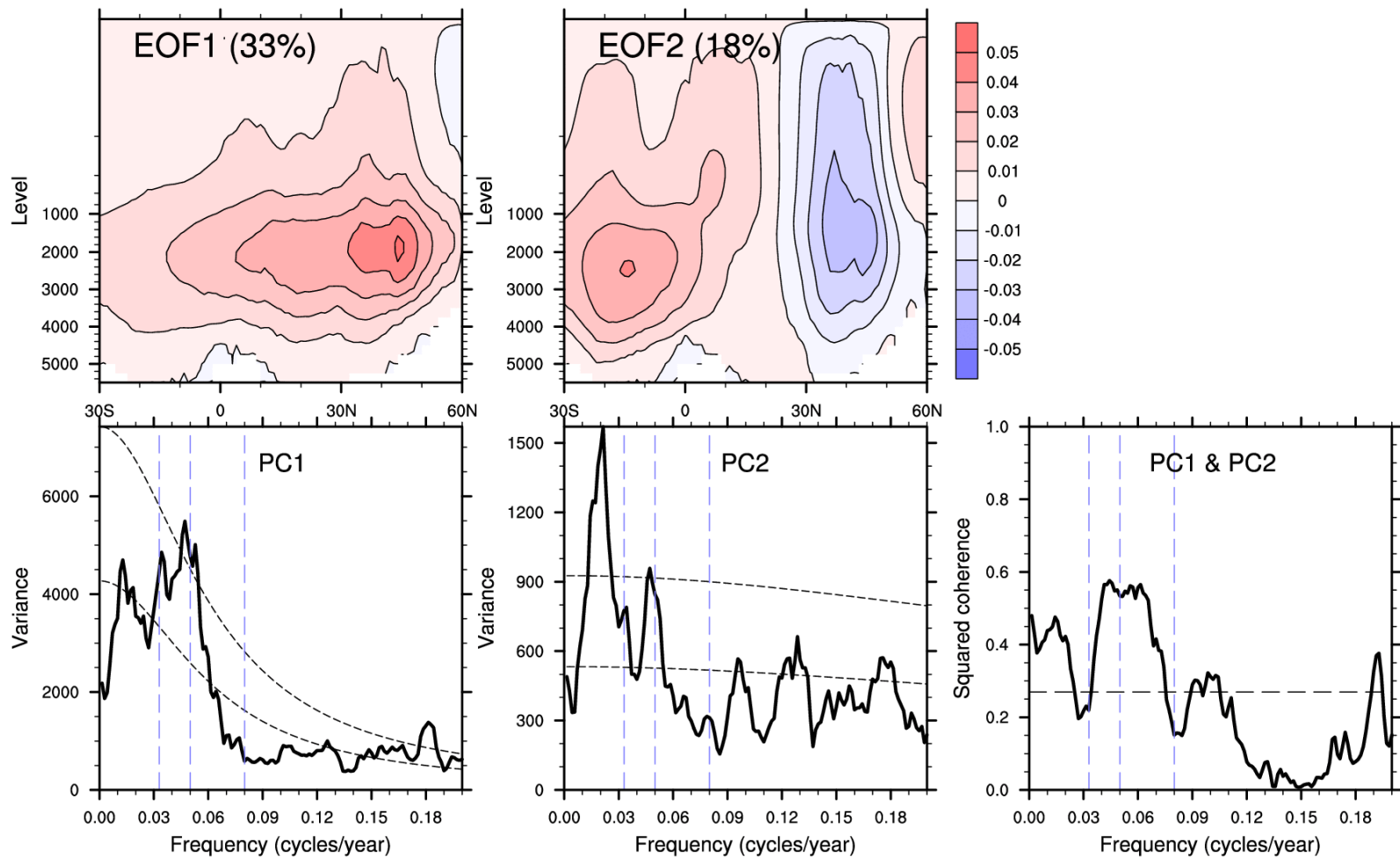
$$\bar{T}_{A1B}(t) = \bar{T}_{1999} + k(t - 1999)$$

$$\bar{T}_{commit}(t) = \bar{T}_{1999} + A(1 - e^{-t/\tau})$$

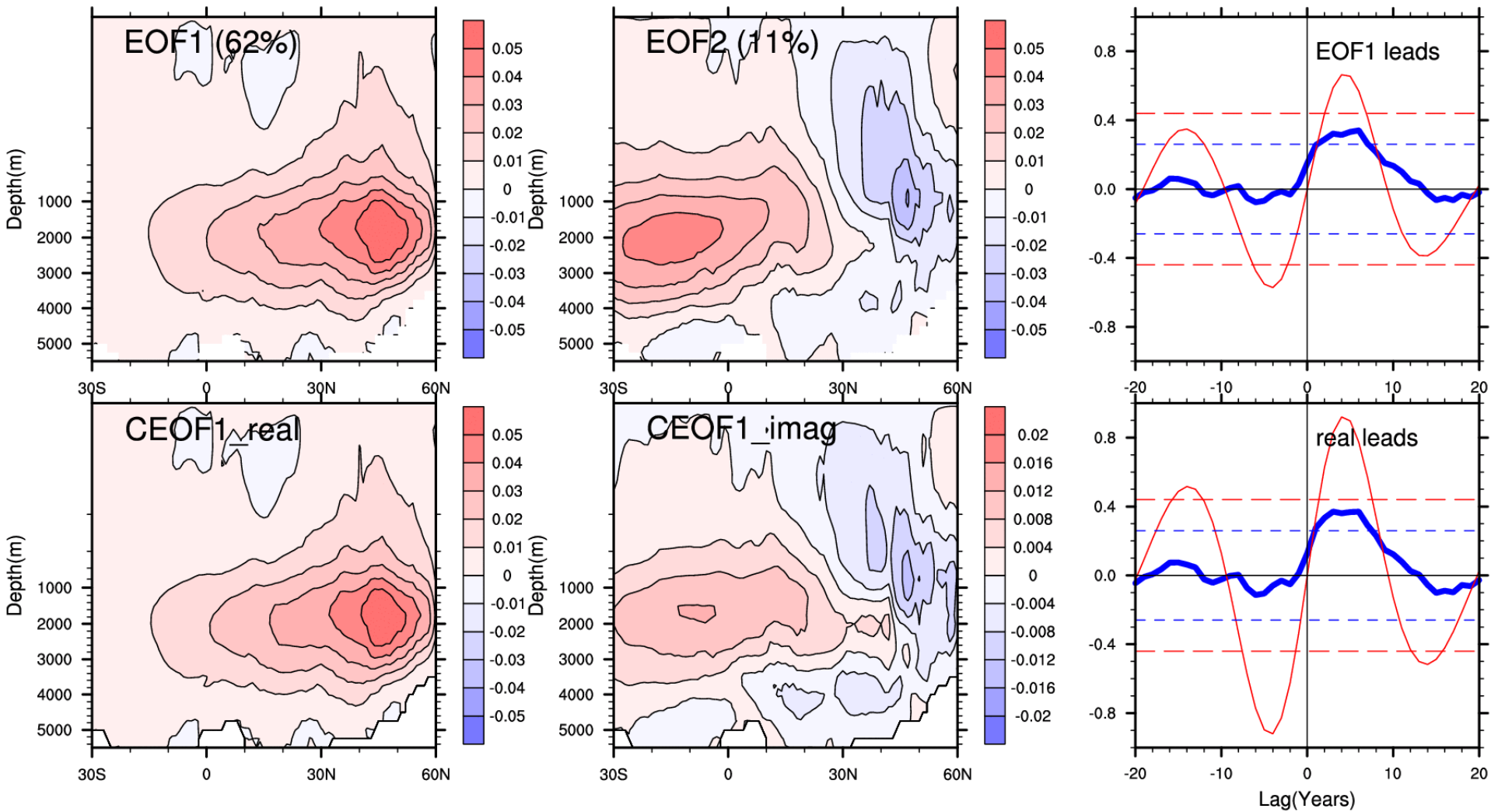
- ✓ Initial-value components
- ✓ Forced response



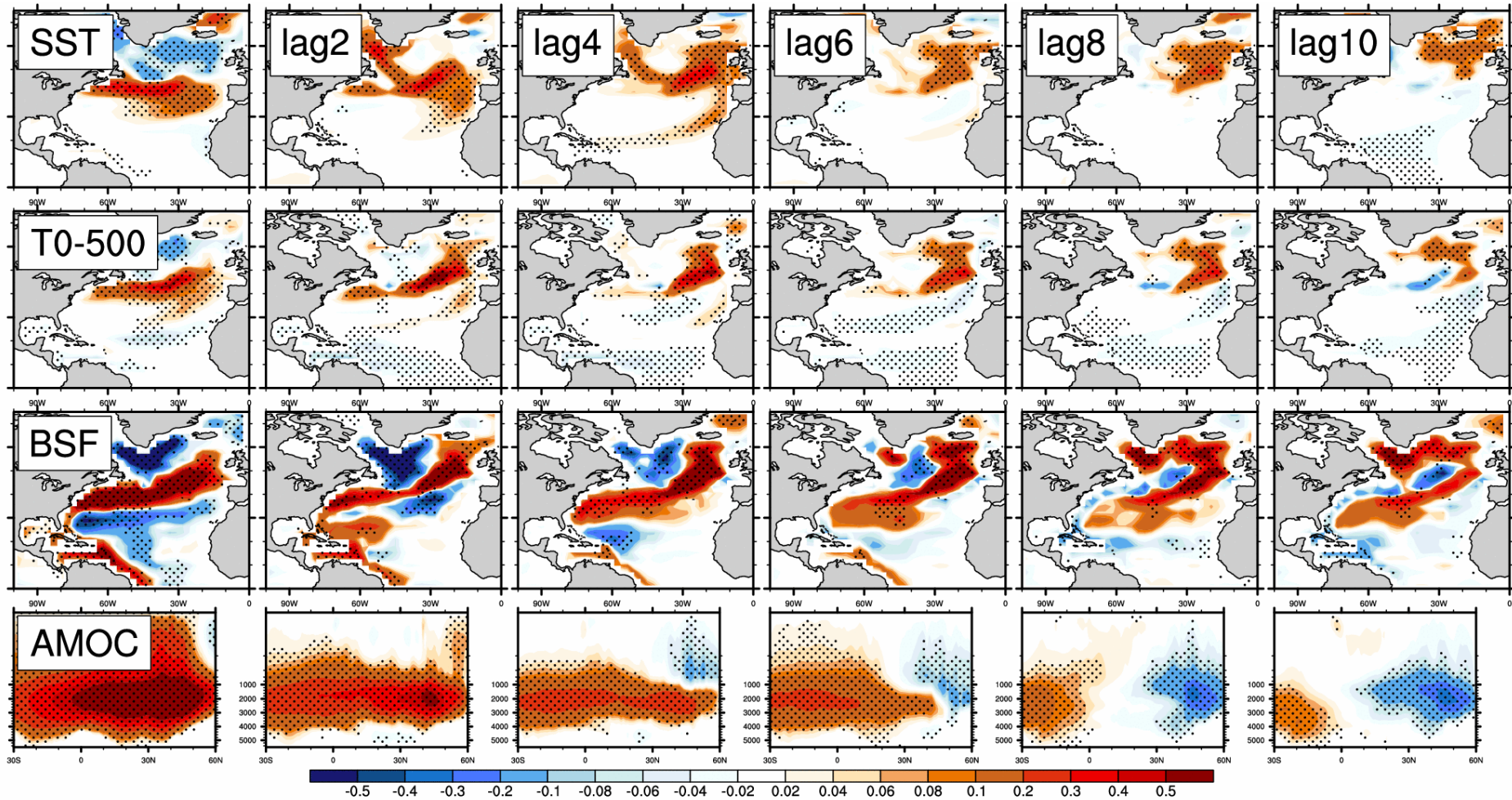
AMOC EOF1 & EOF2 in control run



10-30-yr AMOC EOF1,2 & CEOF1

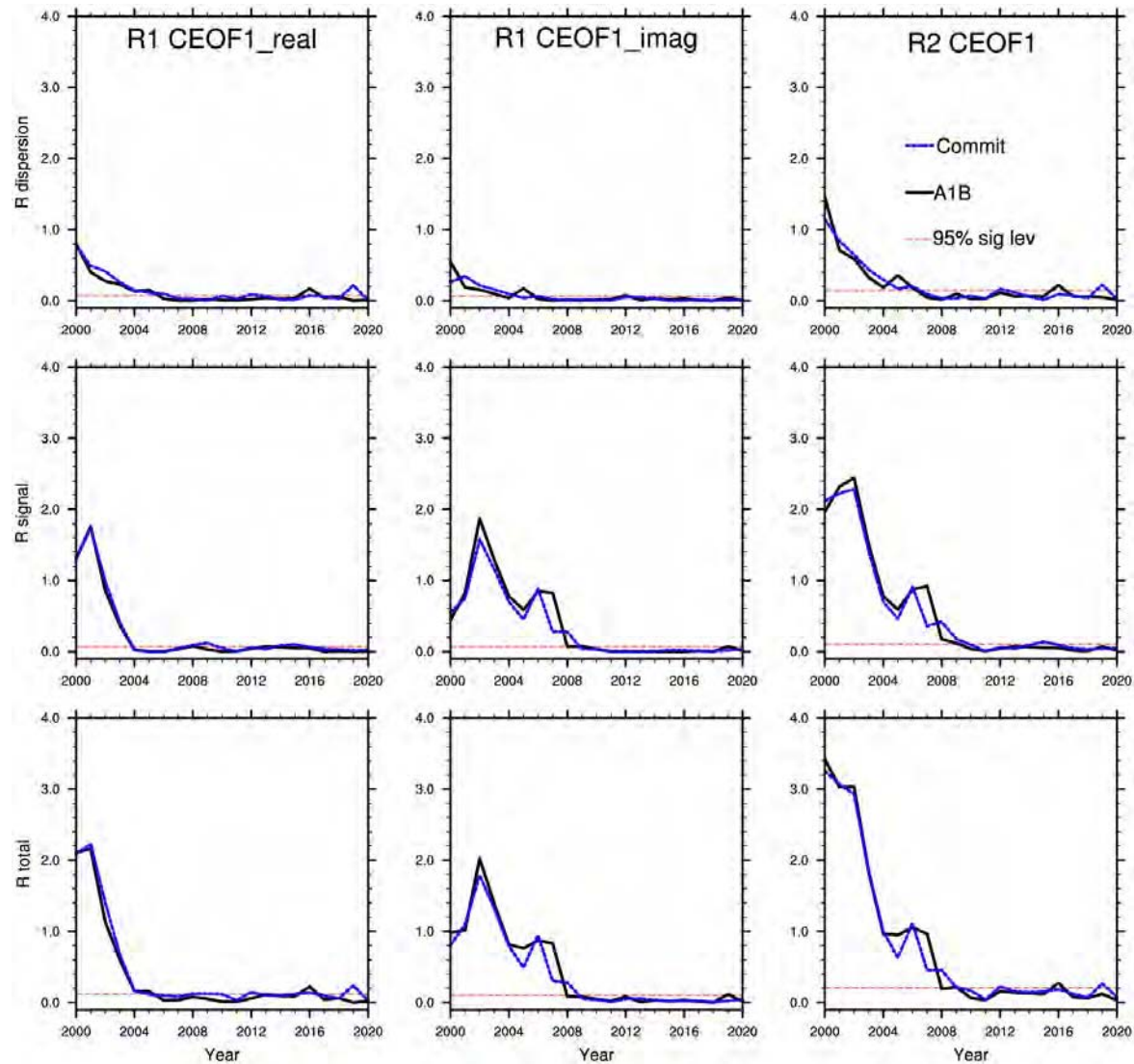


Lag regression to AMOC PC1

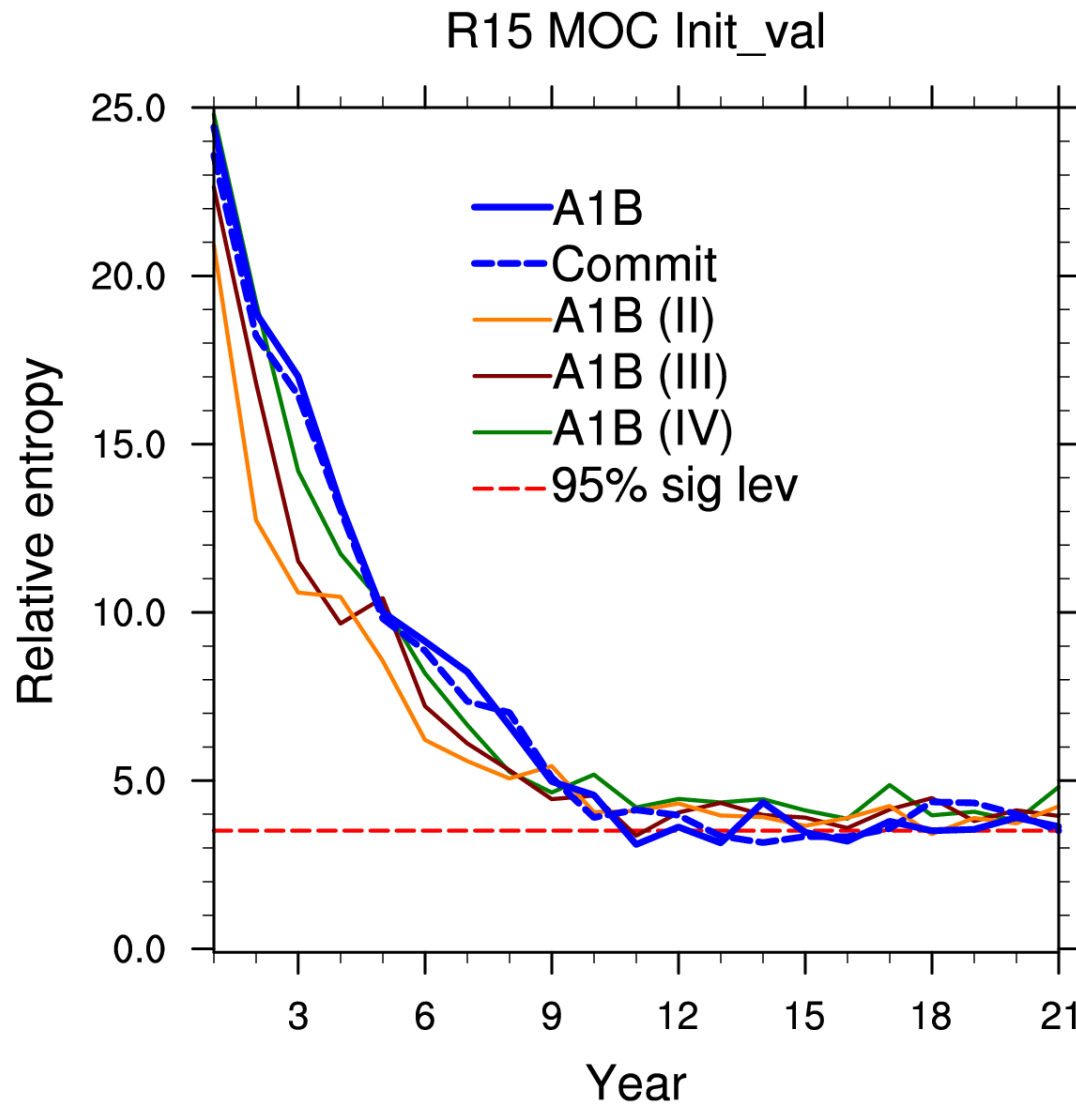


SST 1-2-1 smoothed

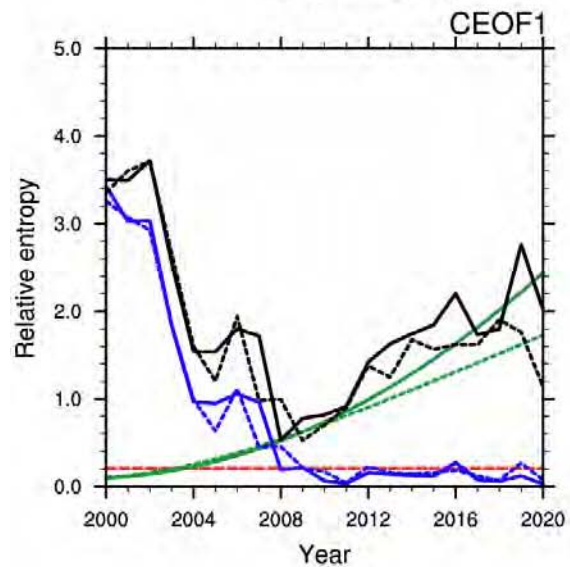
Initial-value Predictability of AMOC CEOF1



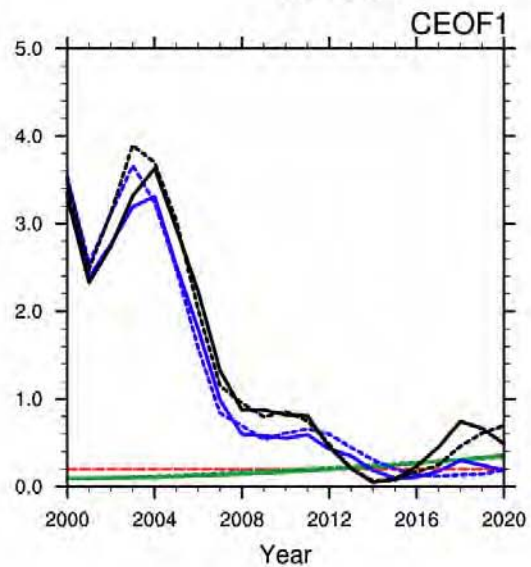
Initial-value Predictability of the AMOC (15 EOFs)



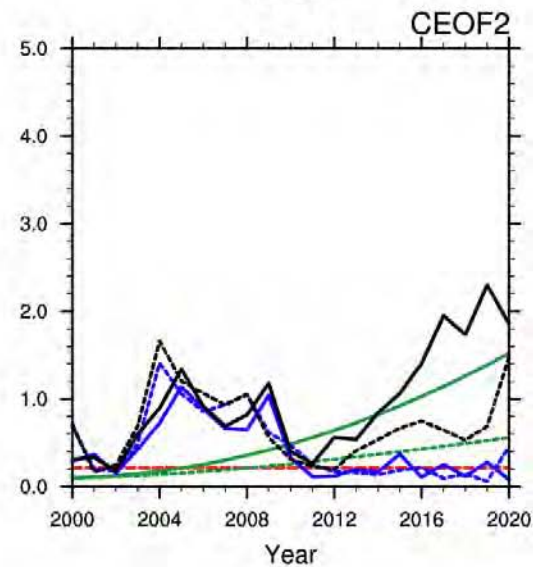
AMOC



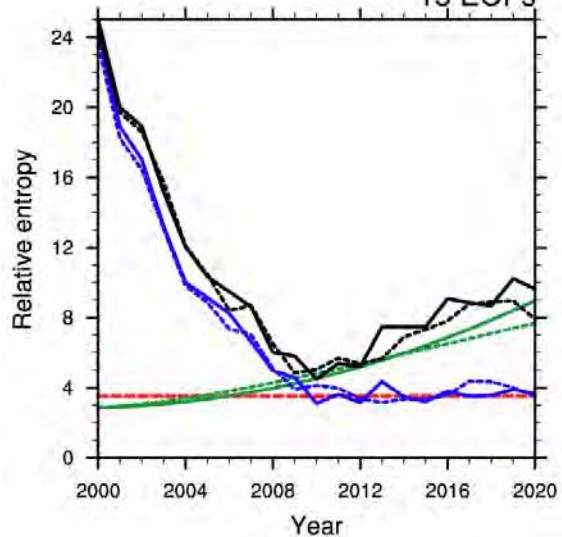
T0-500



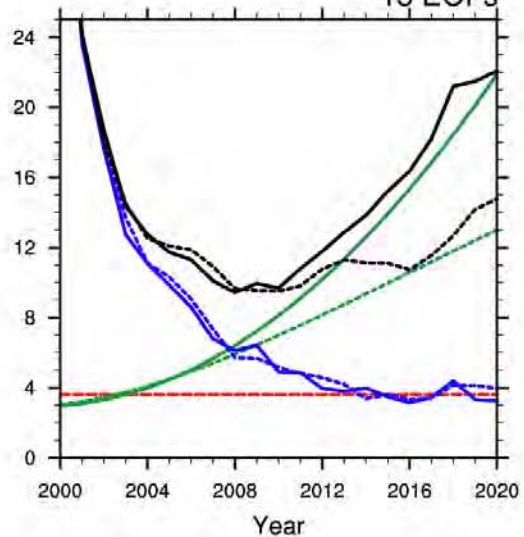
SST



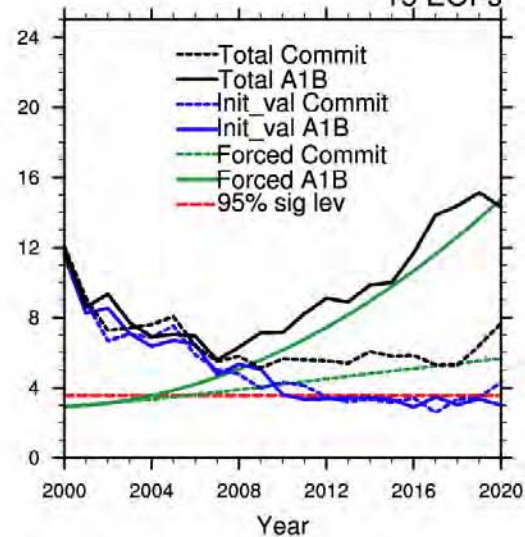
15 EOFs



15 EOFs



15 EOFs



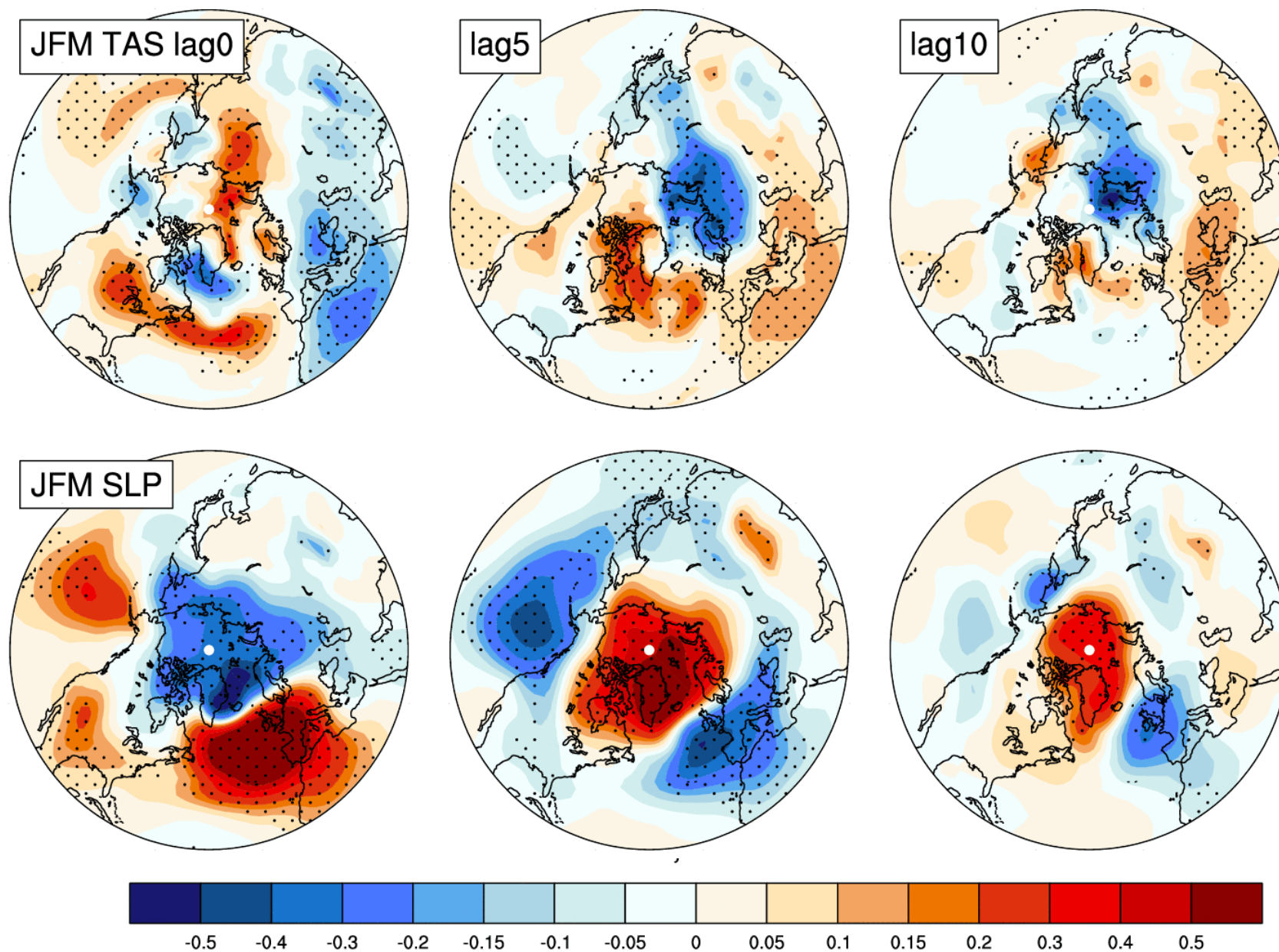
AMOC

- The initial-value component of the AMOC is predictable for about 15 years in the CCSM3 large ensemble experiments.
- The predictability mainly comes from the signal component.
- There is a tendency of the EOF1 to evolve to the EOF2 on the decadal time scale.
- The forced response is more predictable than the initial-value component after 10 years.
- The AMOC is no more predictable than the subsurface temperature.

atm

- Are there signals in the atmospheric large-scale surface climate?
- Are the signals associated with the AMOC?
- Relative importance of the natural variability and the forced response

Lag correlation from the 700-yr control run

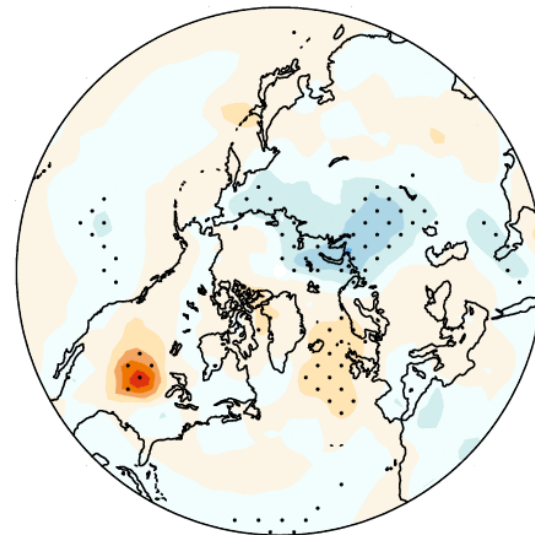
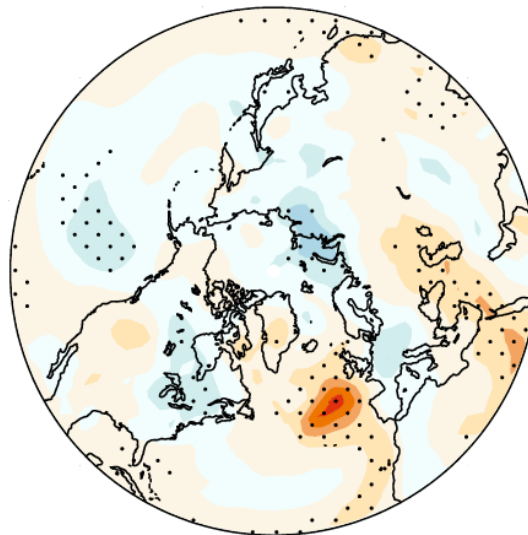
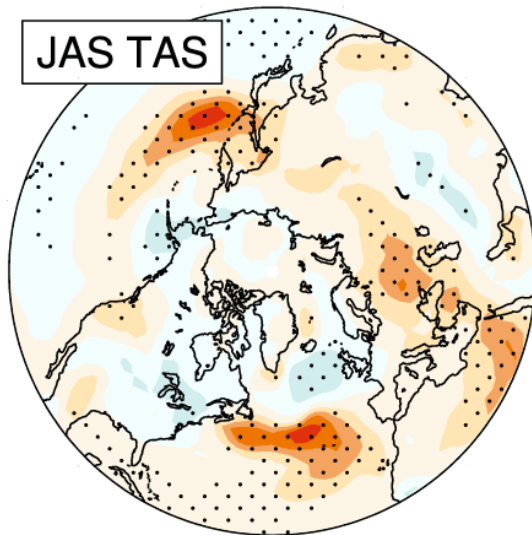


lag0

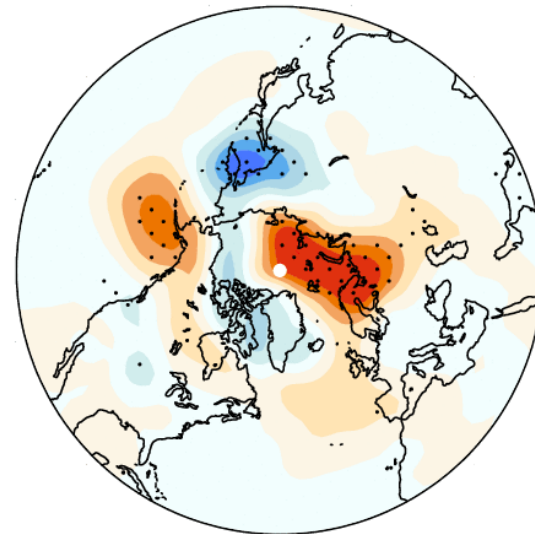
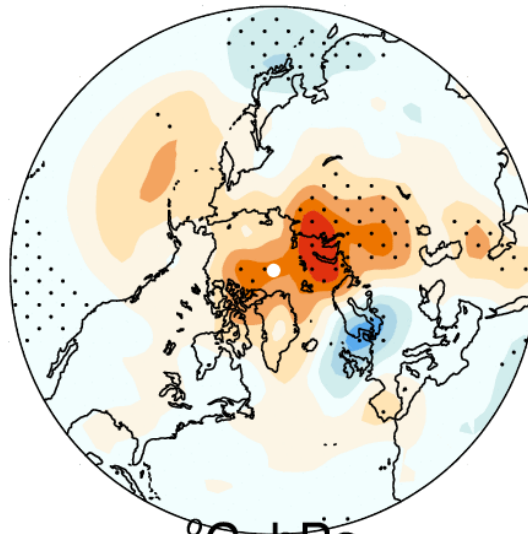
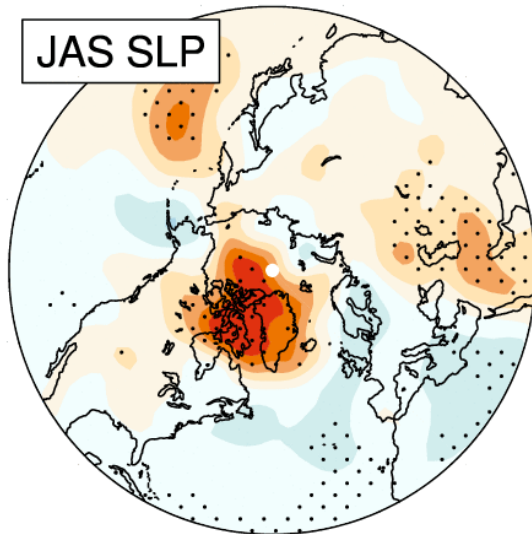
lag5

lag10

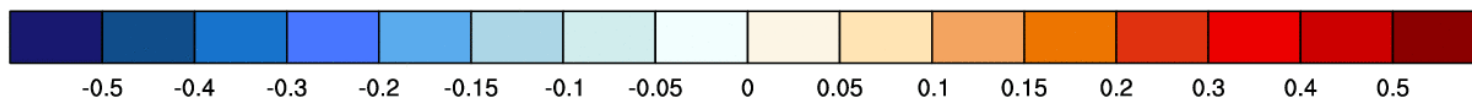
JAS TAS

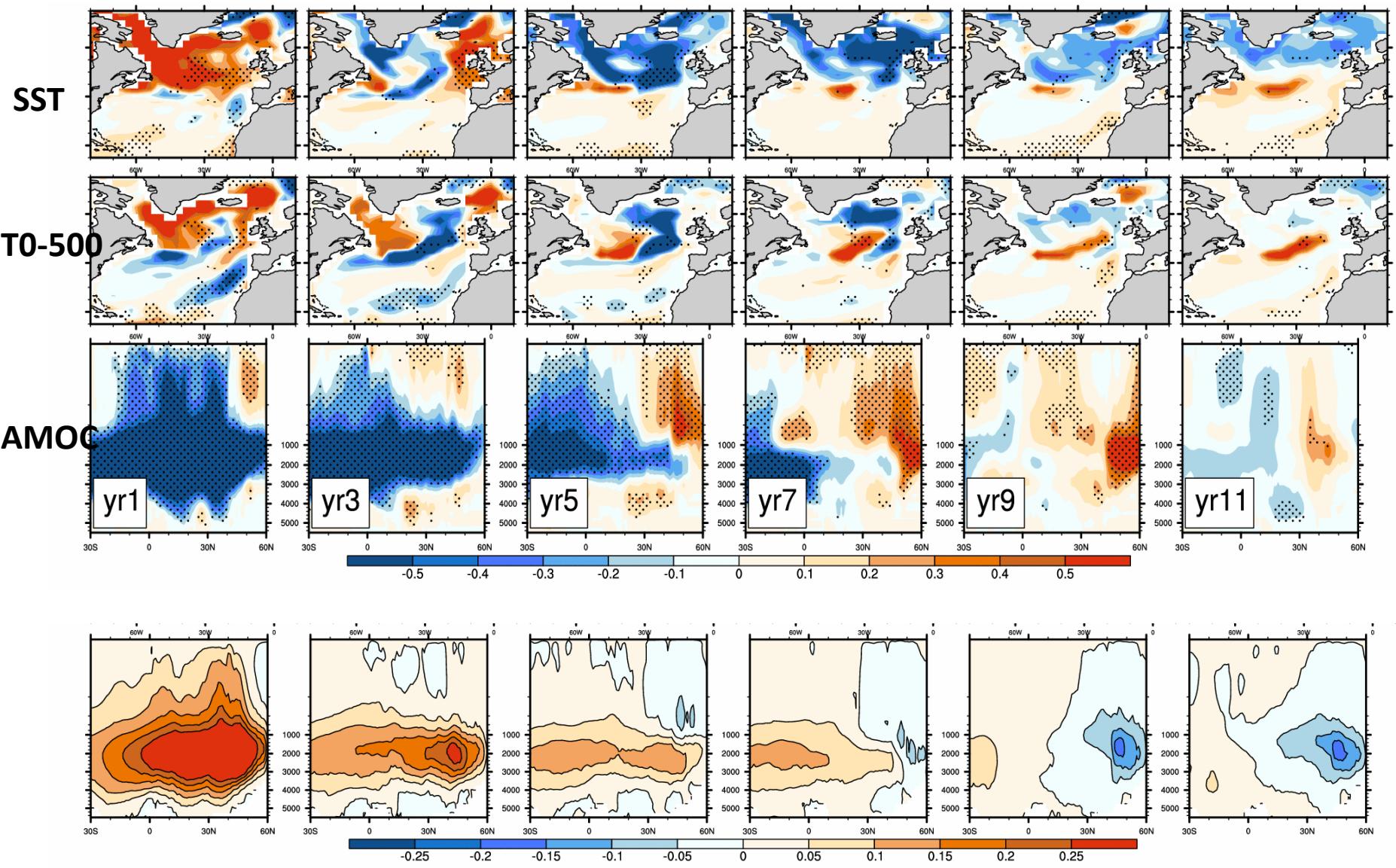


JAS SLP



°C, hPa

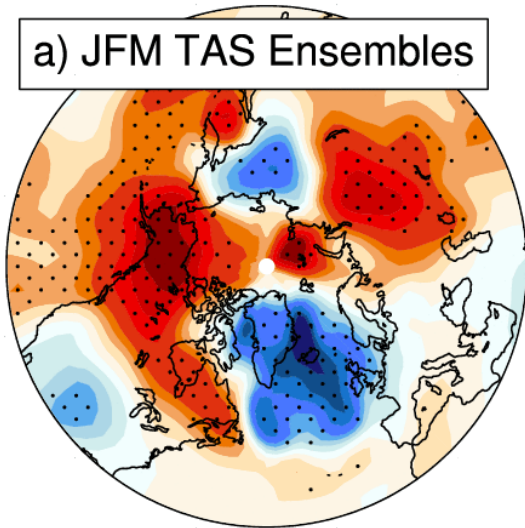




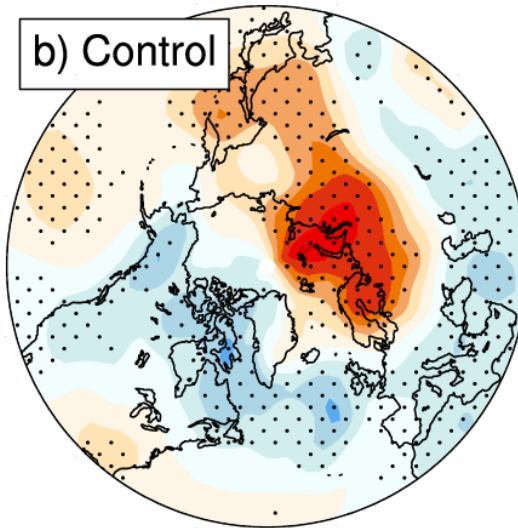
1-2-1 smoothed to remove the ENSO signals

yr5-10

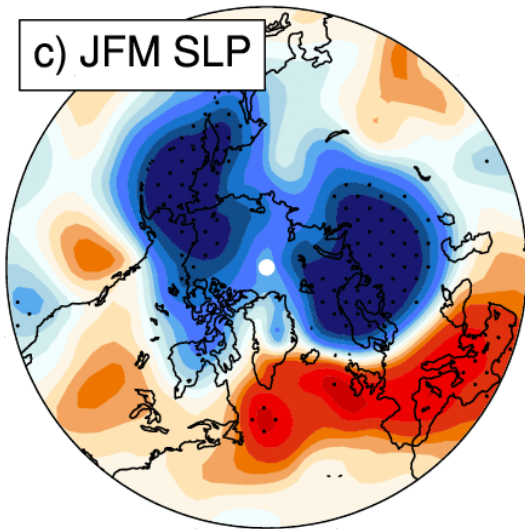
a) JFM TAS Ensembles



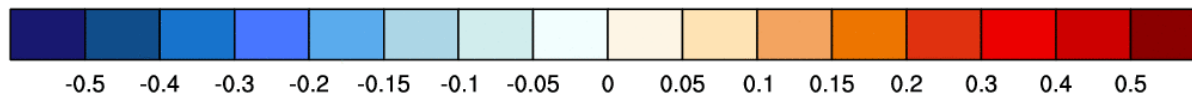
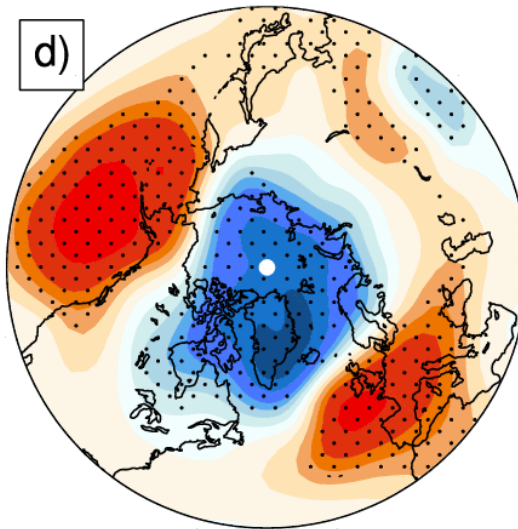
b) Control



c) JFM SLP



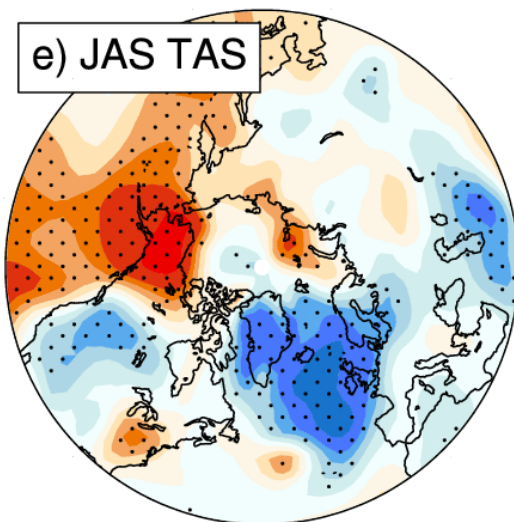
d)



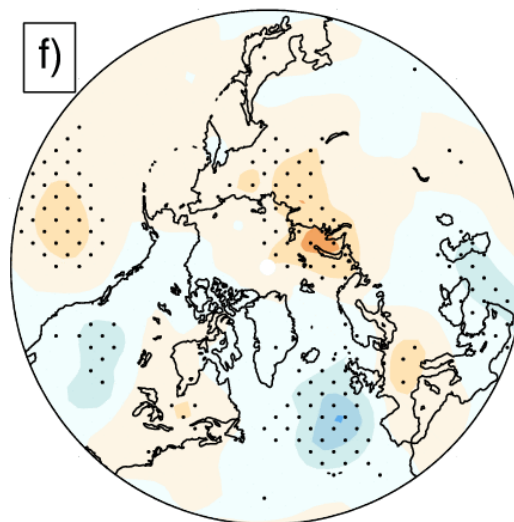
ensemble

control

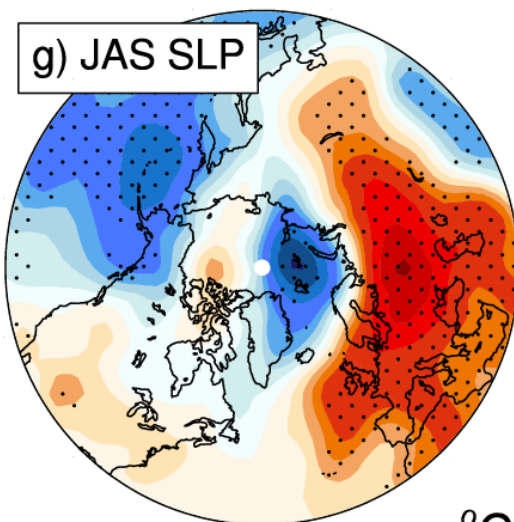
e) JAS TAS



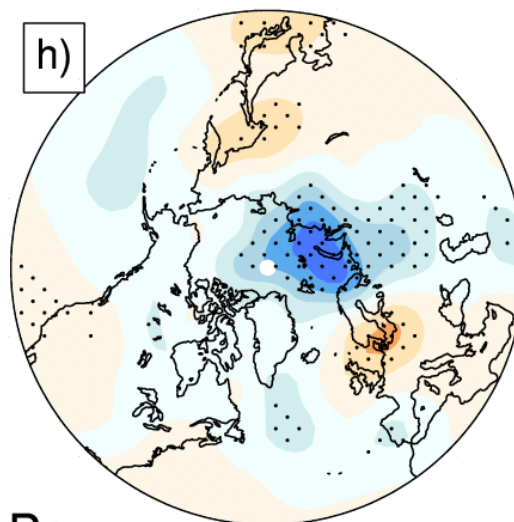
f)



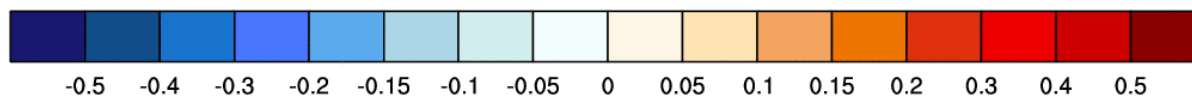
g) JAS SLP



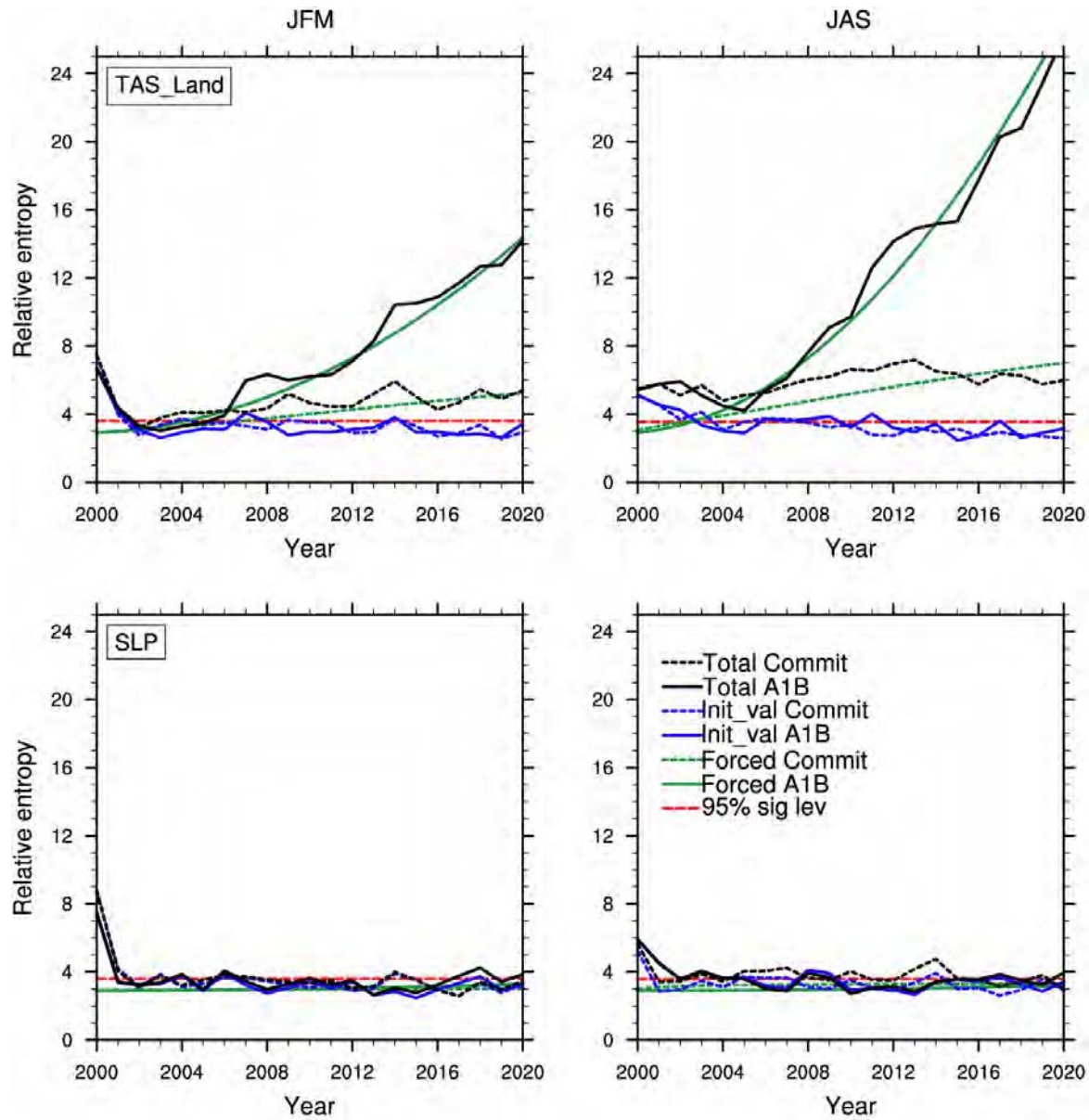
h)



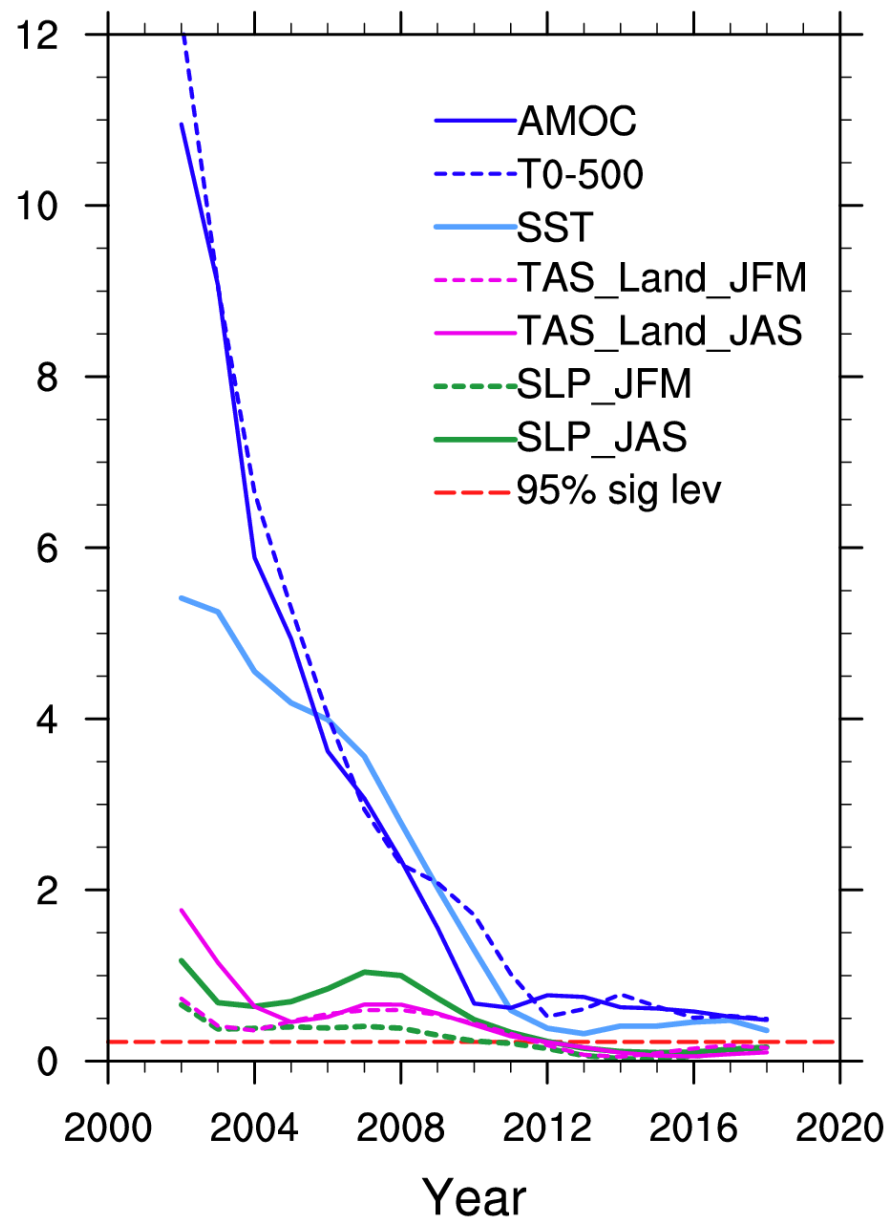
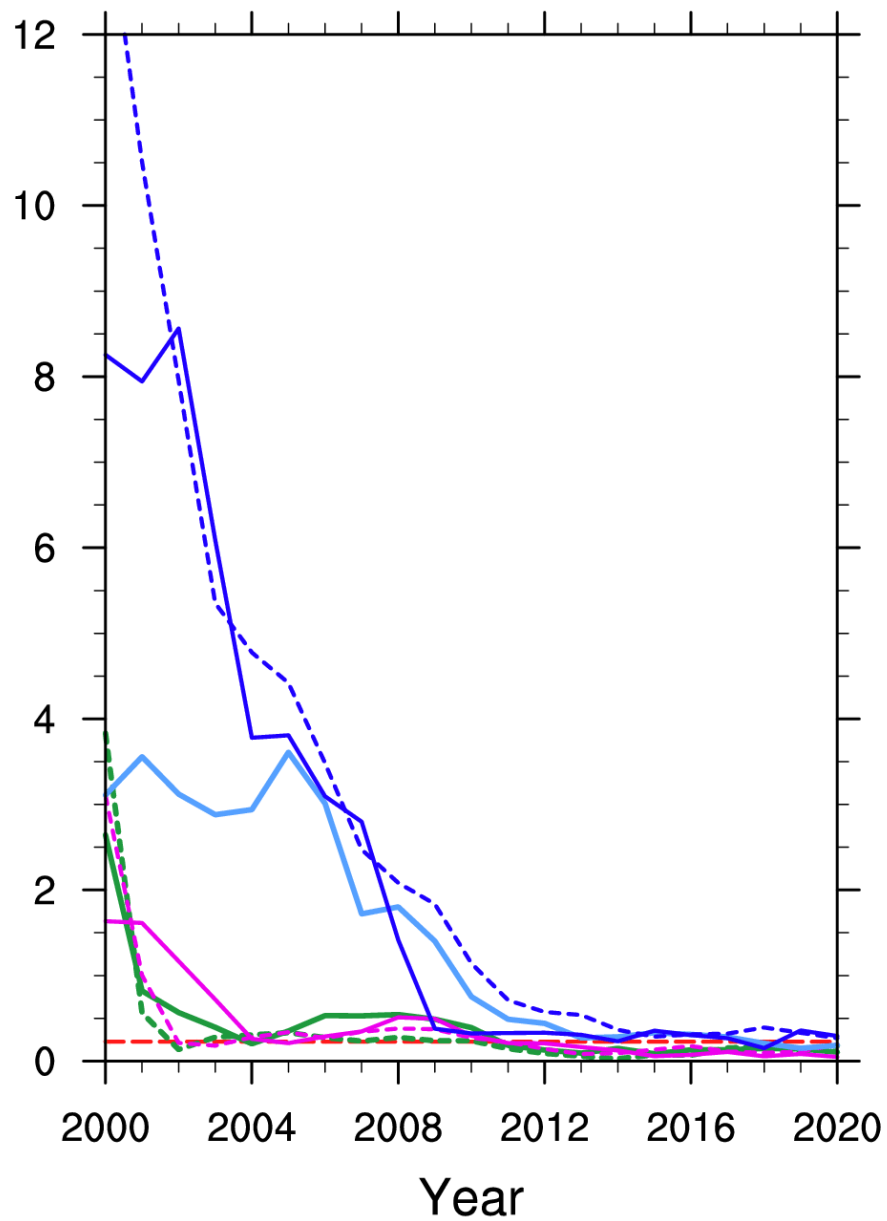
°C, hPa



Relative entropy of 15EOFs in NH (20N-90N)



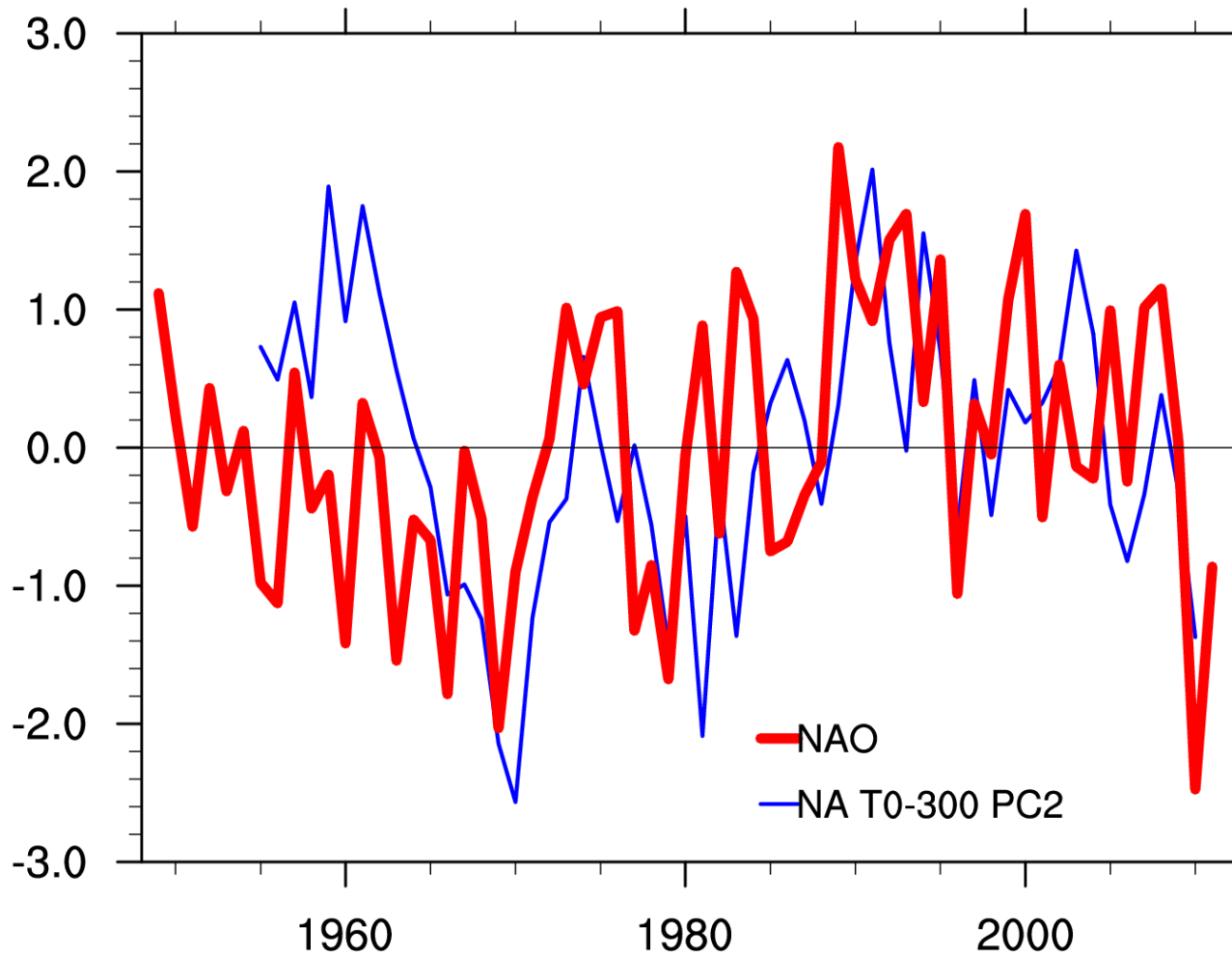
Relative Entropy



Summary of the atm

- The large ensemble mean indicates that the AMOC grows from $-EOF1$ to $-EOF2$ in the first decade. Both the subsurface temperature and SST signals are consistent with the control run composite at the same AMOC phase.
- The subpolar gyre SST anomalies associated with the AMOC variability can influence the atmosphere and produce surface climate predictability that goes beyond the ENSO time scale. However, the resulting initial-value predictability in the atmosphere is very weak.

Multidecadal Variability in the NAtl .



5-yr running mean

Levitus upper300m T

