

What can we say about the coming decade?

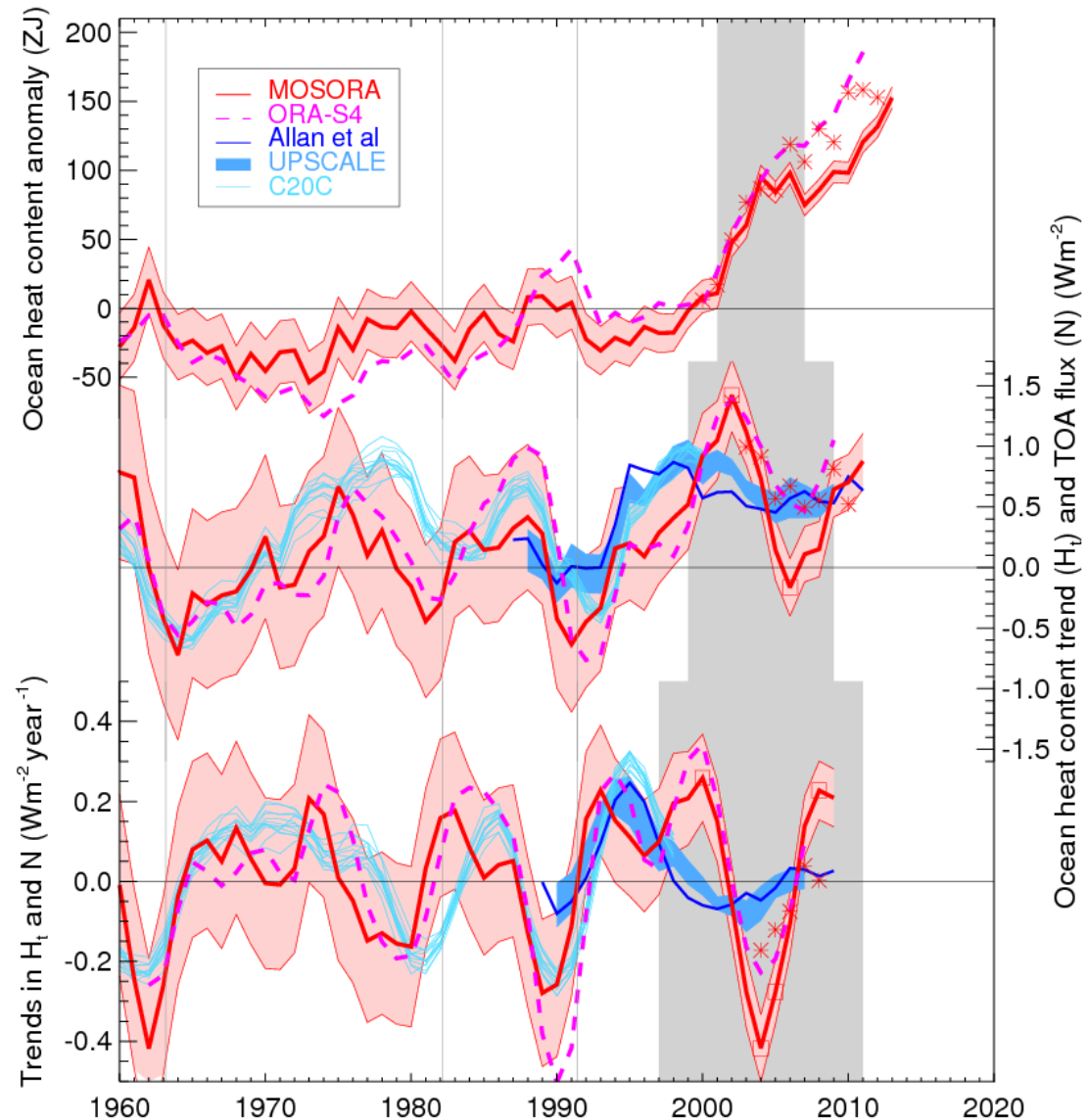
Doug Smith, Nick Dunstone, Rosie Eade, Leon Hermanson, Adam Scaife, ...



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Hadley Centre

- Good agreement ($r=0.8$) between Met Office (MOSORA) and ECMWF (ORA-S4) H_t
- Reasonable agreement ($r=0.6$) between H_t and N from 1960 to 1999
- Also seen in gradients (N_t vs H_{tt})

Observed energy budget

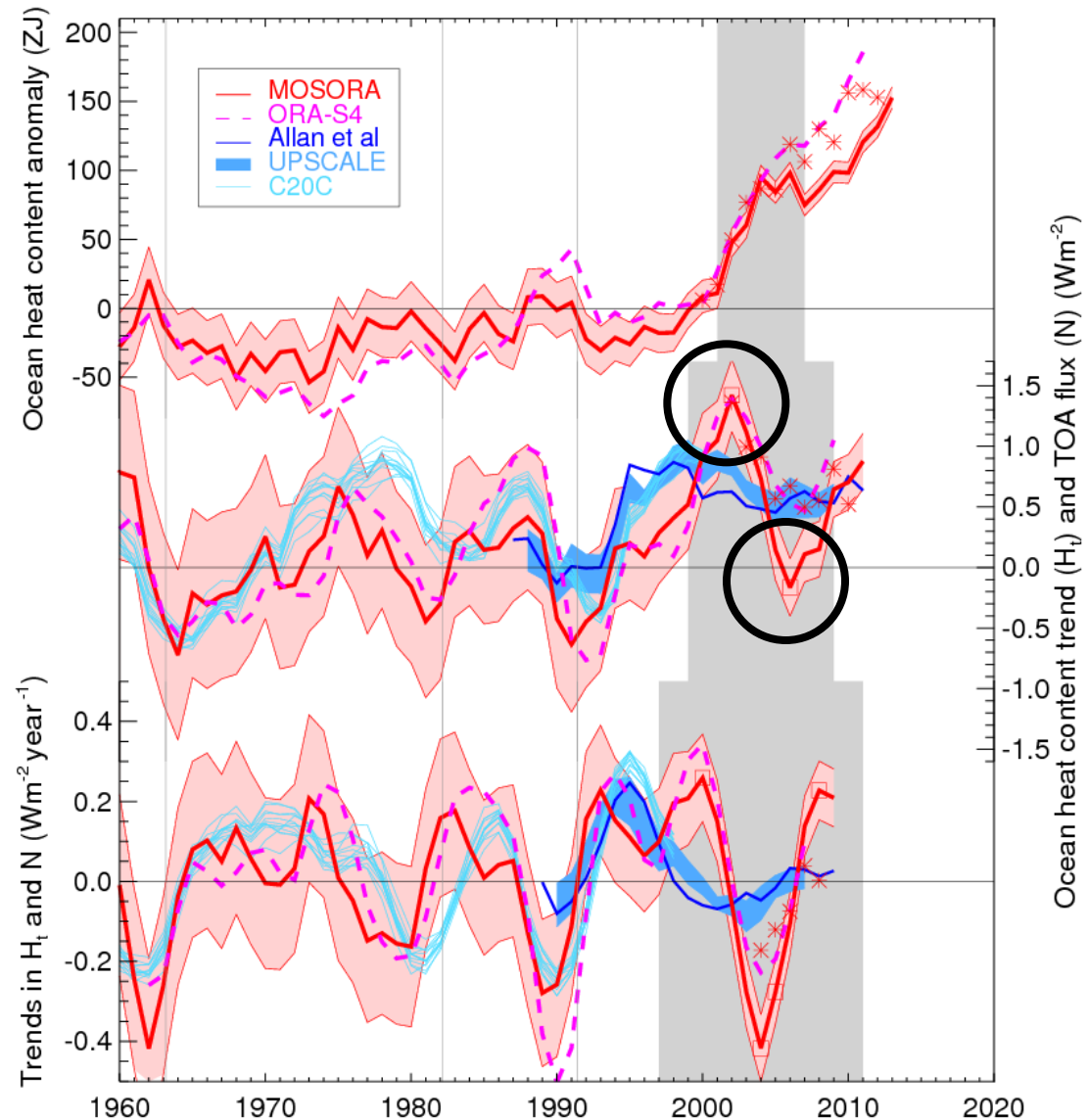




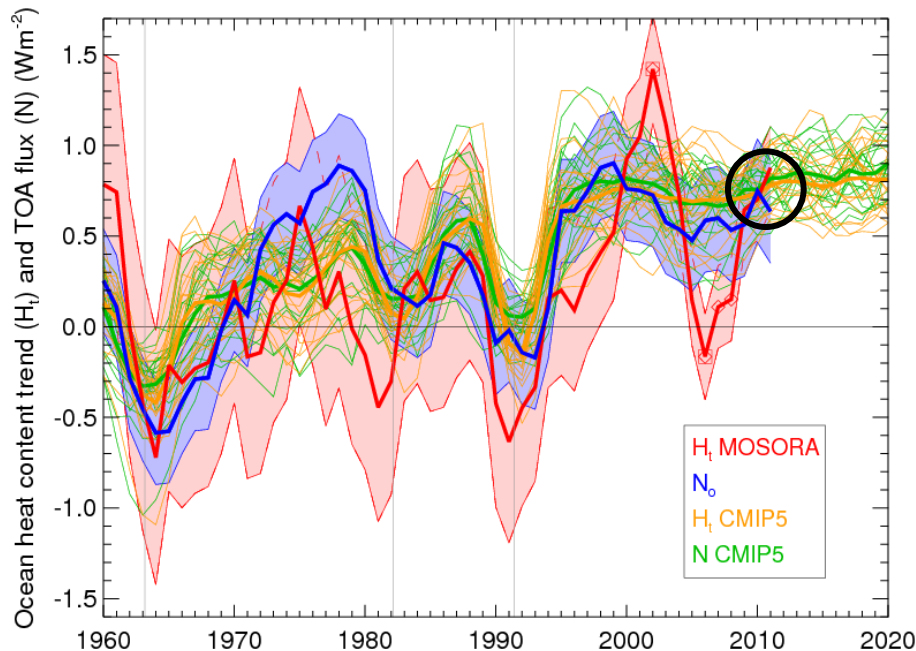
Met Office
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Observed energy budget

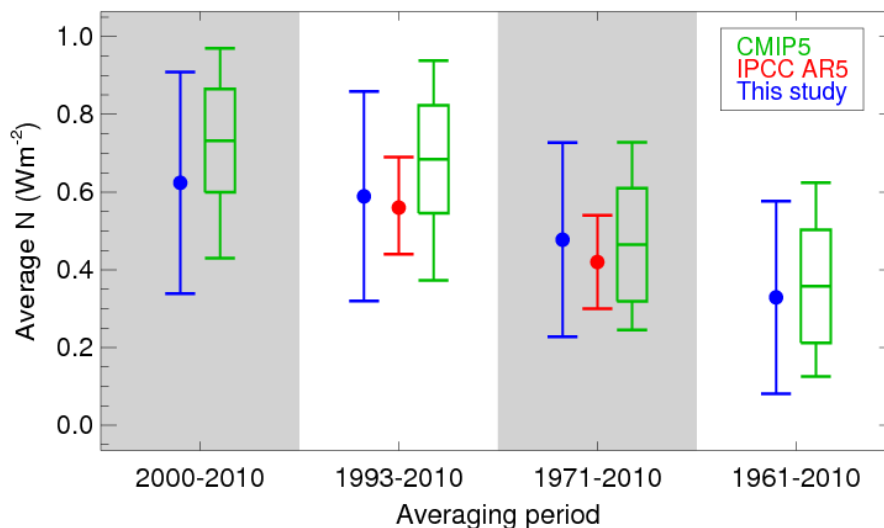
- Maximum H_t in 2002 is inconsistent with N from satellites and AMIP simulations, so likely to be spurious
- Ocean analyses between 2000 and 2007 (grey shading) may be unreliable
- Also seen in analyses without Argo (red asterisks)



Comparison with CMIP5 coupled models

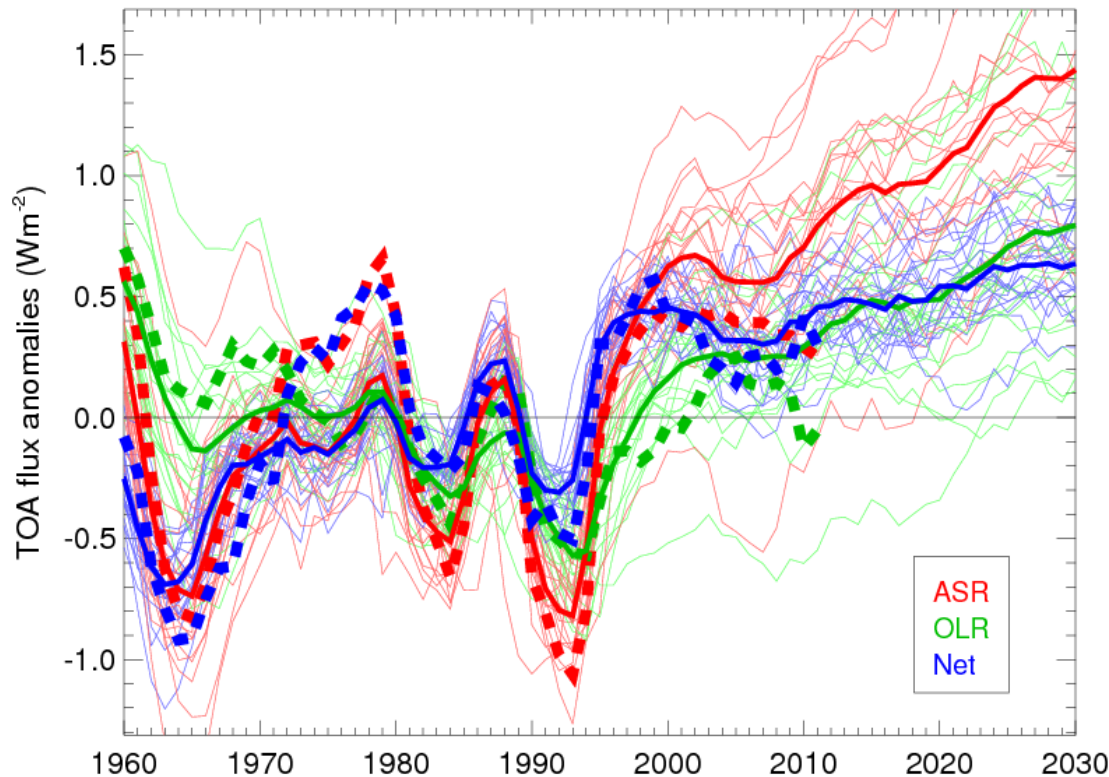


- Observation-based estimate N_o created by averaging satellite and AMIP simulations
- Goes back to 1960
- Compare with CMIP5 coupled models
 - Need to remove CMIP5 control run values because of drifts
- Good agreement ($r=0.82$) with ensemble mean of CMIP5 coupled models
- Variations dominated by volcanic eruptions (Agung 1963, El Chichon 1982 and Pinatubo 1991)



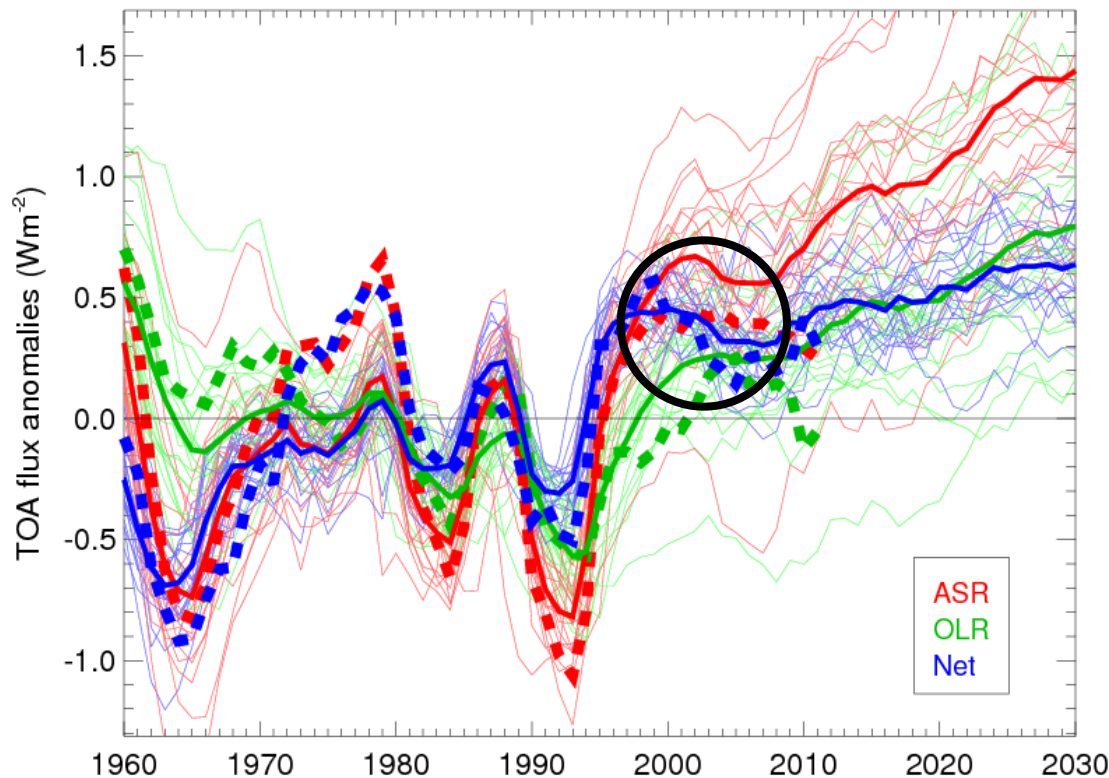
- Absolute values of N for different periods in reasonable agreement with CMIP5 and IPCC
- BUT large uncertainties, dominated by estimates of H_t needed to anchor N – note disagreement in changes in N and U even in the most recent estimates

Observation-based and model anomalies



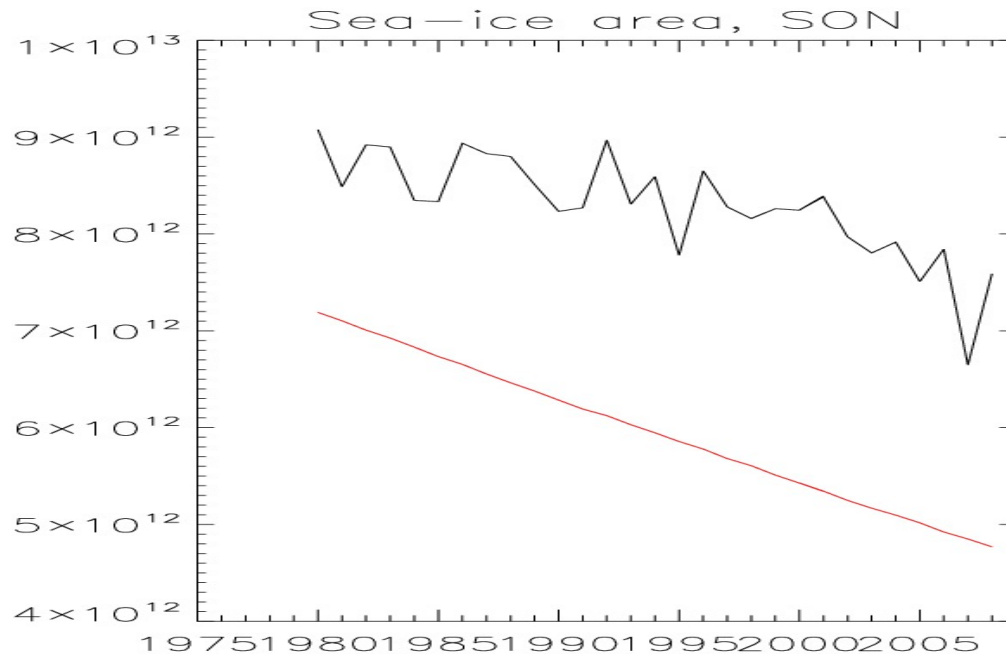
- Anomalies wrt 1960-2011 → overcomes anchoring uncertainties
- Radiative components: absorbed shortwave (ASR, red), outgoing longwave (OLR, green)
- $N = ASR - OLR$
- Good agreement between models (solid) and obs (dashed) for ASR (red, $r=0.87$) and OLR (green, $r=0.80$)
- Models simulate response to volcanoes reasonably well (at least in ensemble mean)

Observation-based and model anomalies



- Observation-based decline in N of $0.31 \pm 0.21 \text{ Wm}^{-2}$ between late 1990s and mid-2000s
- Spans different satellites – but also seen in AMIP simulations
- Role in warming pause?
- Consistent with minor volcanoes, reduced solar activity
- But driven by increasing OLR rather than reduced ASR? – though ASR lower than models
- N may also change through internal variability
- Hence relative roles of external forcing and internal variability remain unclear
- Increased uptake of heat by the ocean might not be crucial for explaining the hiatus

Role of Arctic sea ice?



- 30 year period 1979 to 2009
- Control: use observed Arctic sea ice
- Perturbed: use reduced Arctic sea ice, obtained by projecting the 30 year trend
- 10 ensemble members (300 simulated years)
- AMIP: atmosphere model forced by observed SST (set to 0°C where sea ice removed)
- COUPLED: fully coupled model, constrained by observed temperature and salinity below 200m – **SST can respond to sea ice**

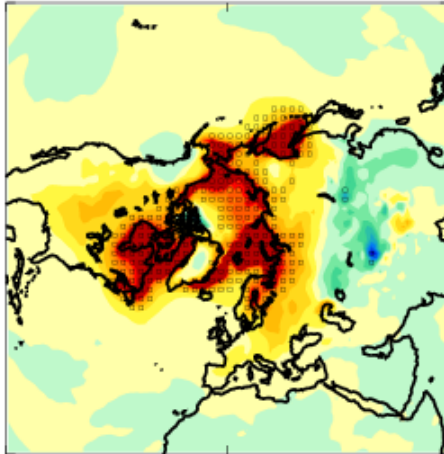


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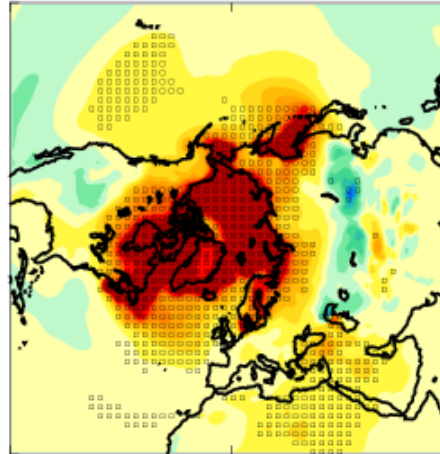
Temperature

DJF

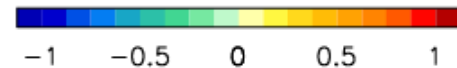
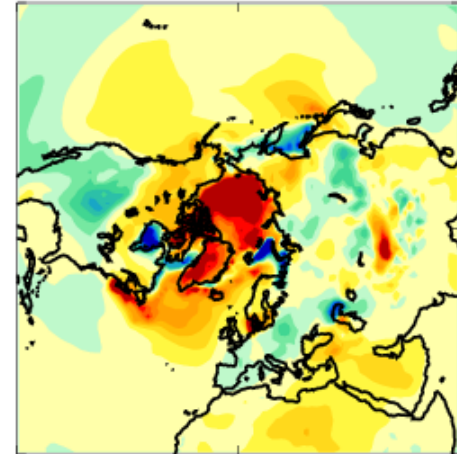
AMIP



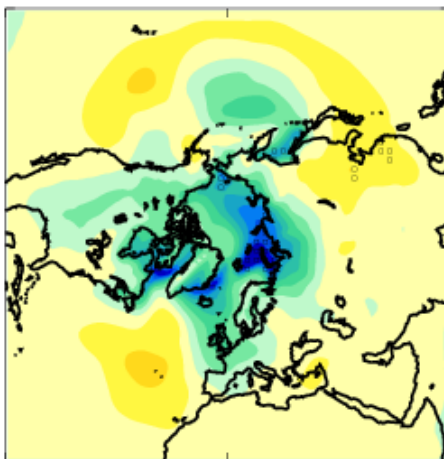
Coupled



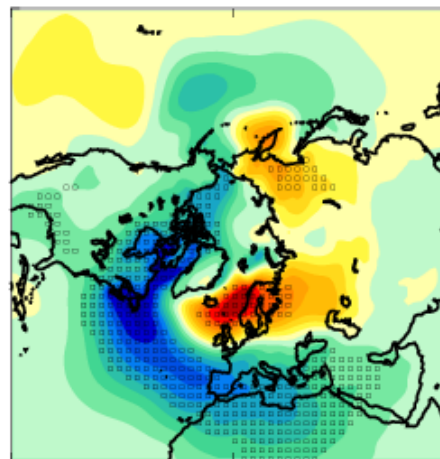
Difference



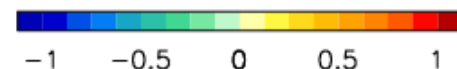
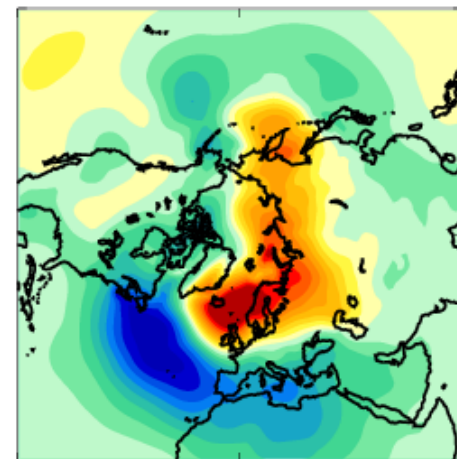
AMIP



Coupled

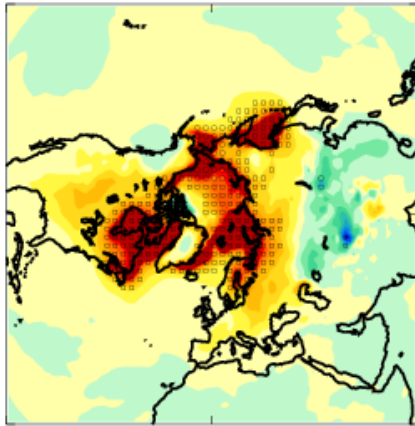


Coupled-AMIP

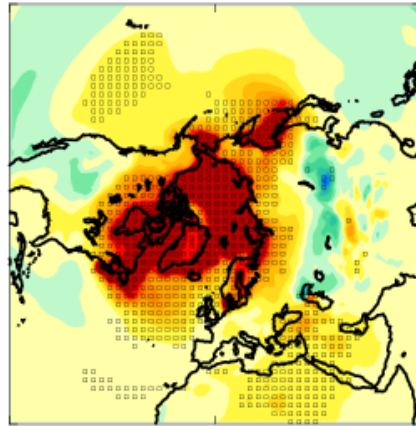


Why is the DJF response different in AMIP and Coupled?

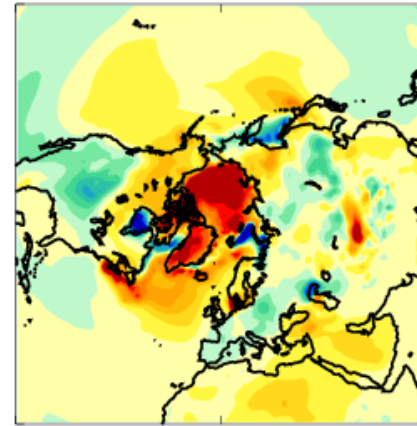
AMIP



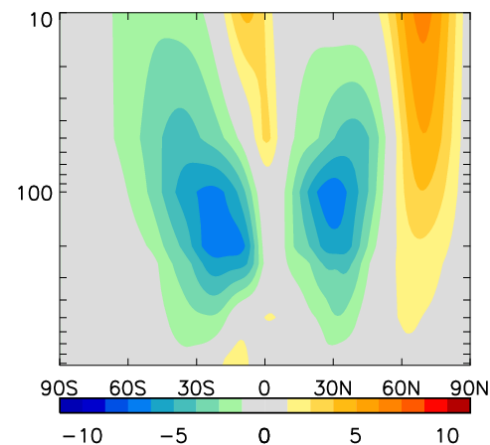
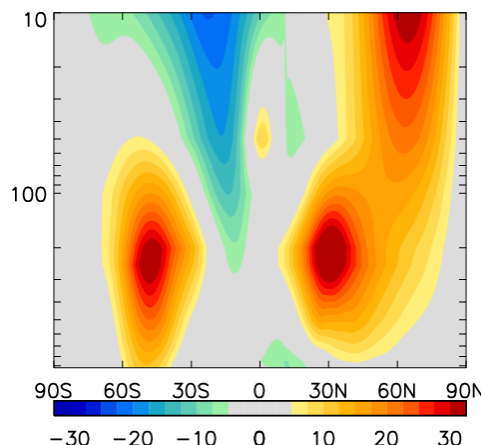
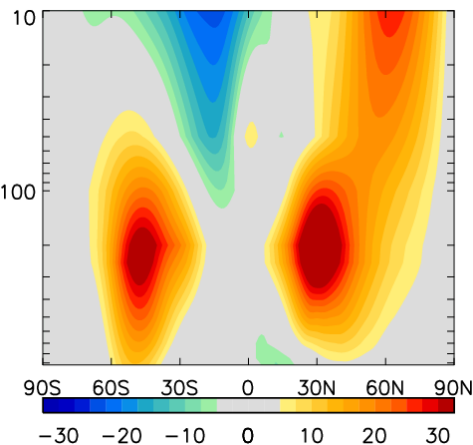
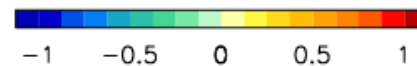
Coupled



Coupled - AMIP

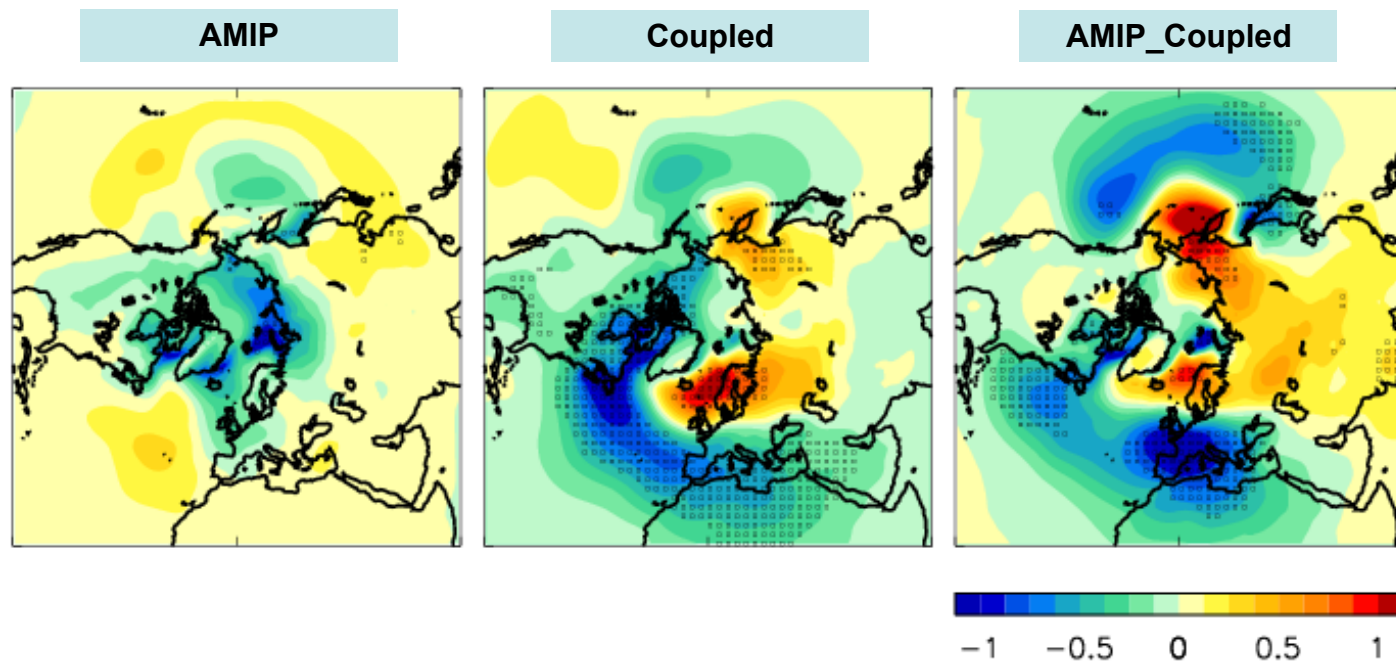


Differences in temperatures: coupled model allows surrounding regions to warm



Model biases: zonal winds are weaker and shifted northwards in the coupled model

DJF sea level pressure

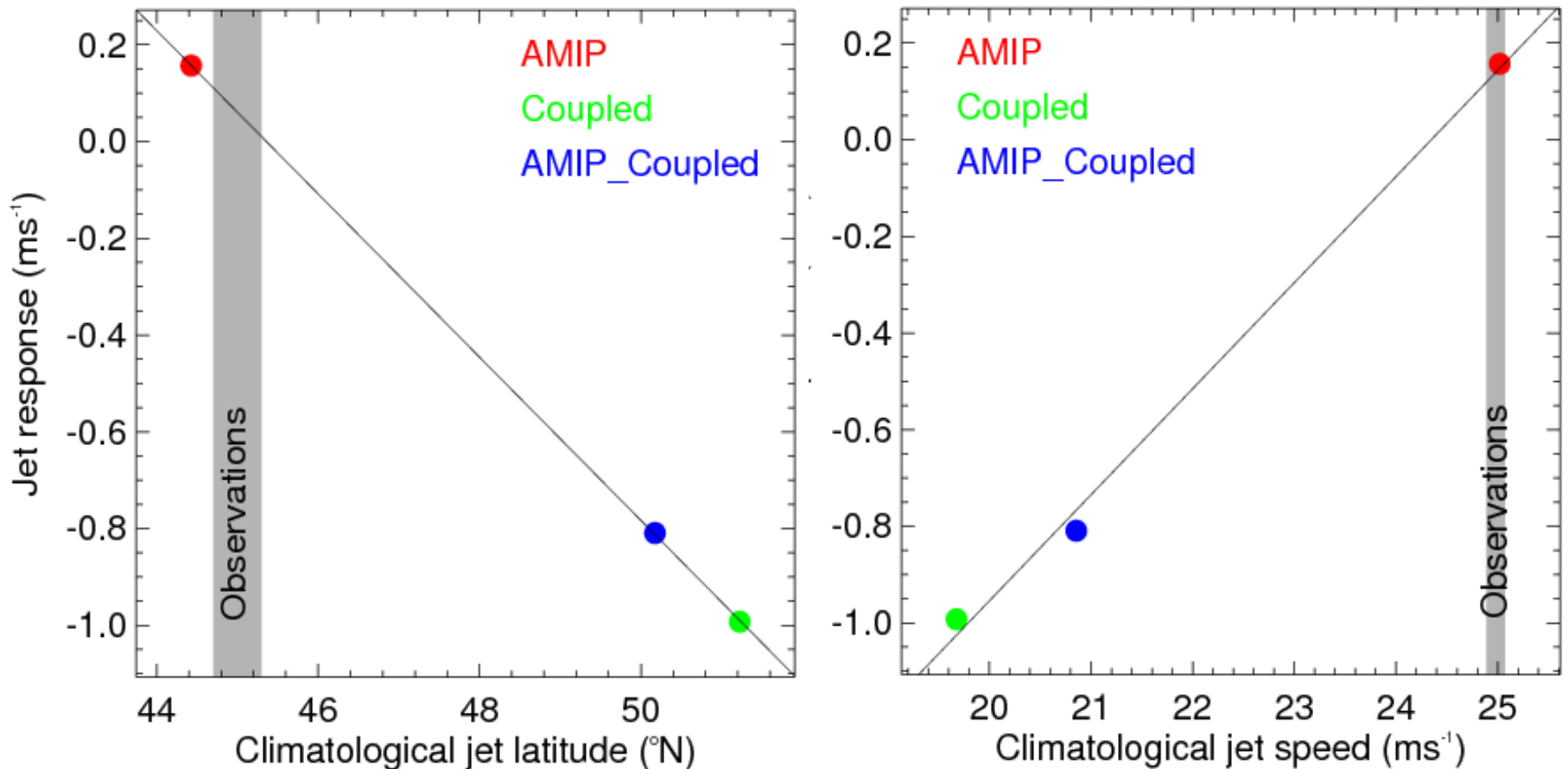


- **AMIP_Coupled: as AMIP but with coupled model SST biases**
- **AMIP_Coupled reproduces the Coupled model response**
- **Response depends on model climatology**



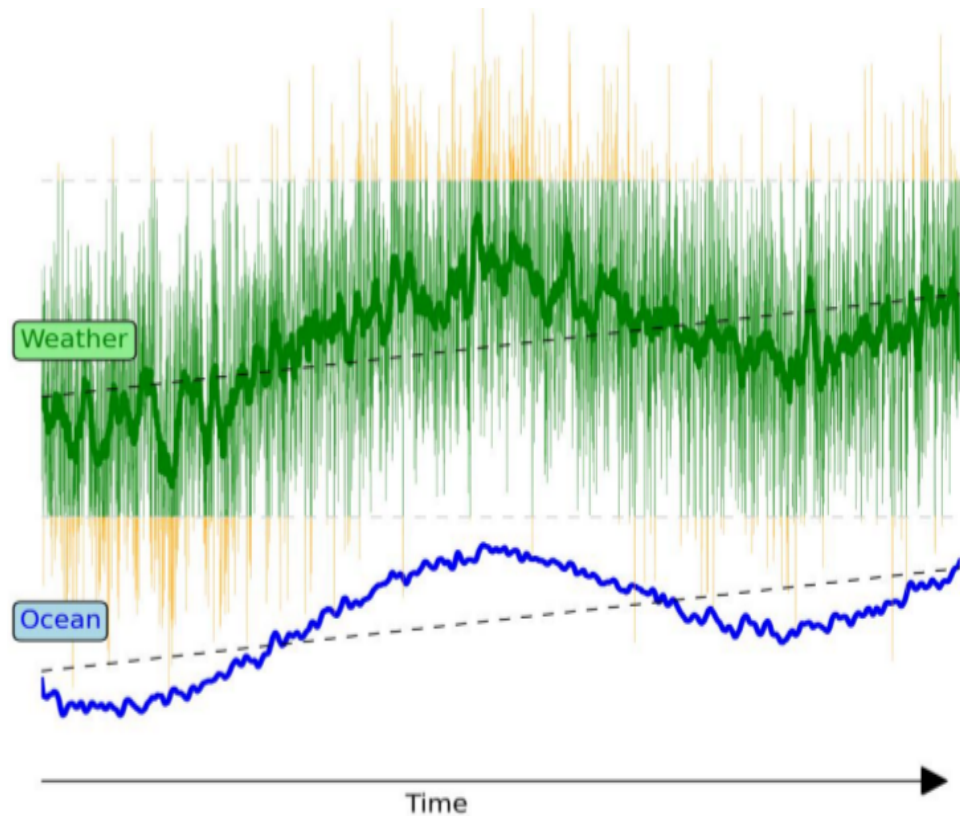
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Jet response vs climatology



- Response depends on the climatological latitude and/or speed of jet
- WHY??

Predictable component



Total variance of a time series can be split into

- Predictable component (*signal*): fraction of variance that is predictable
- Unpredictable component (*noise*)
- Predictable component arises from e.g. ocean influence on atmosphere, external forcings ...
- Modulates probability of extreme events

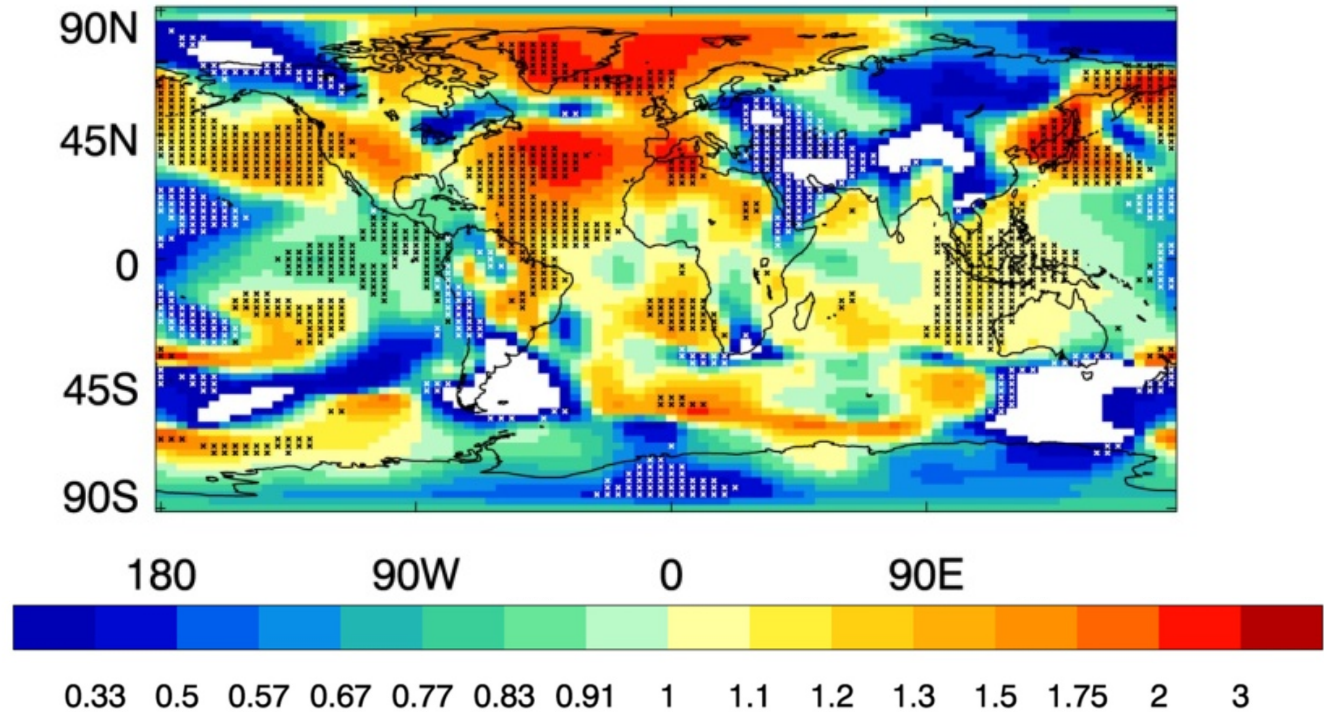
Ratio of predictable components (RPC)

$$RPC = \frac{PC_{obs}}{PC_{mod}} \geq \frac{r}{\sqrt{\sigma_{ens_mean}^2 / \sigma_{ens_members}^2}}$$

- Ignore negative r
- Investigate seasonal in GloSea5; DJF from ~Nov 1st
 - 24 members; 20 start dates (~1st Nov 1992-2011)
- Also decadal (years 2-5), Met Office + CMIP5
 - 70 members; 46 start dates (1st Nov 1960-2005)

RPC in seasonal DJF hindcasts

(b) MSLP



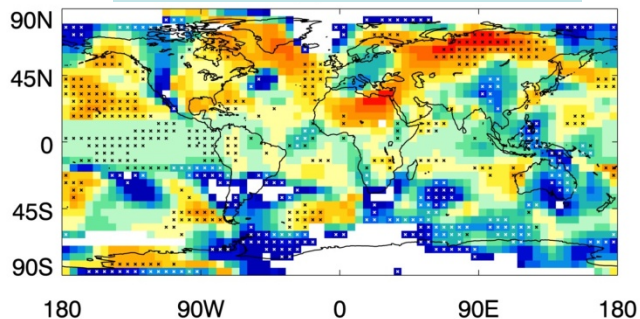
Perfect forecast system: $RPC = 1$

$RPC \ll 1$: over-confidence, ensemble members agree well with one another but not with observations (noted in previous studies)

$RPC \gg 1$: under-confidence, ensemble mean agrees well with observations but members agree less well with each other (unexpected!)

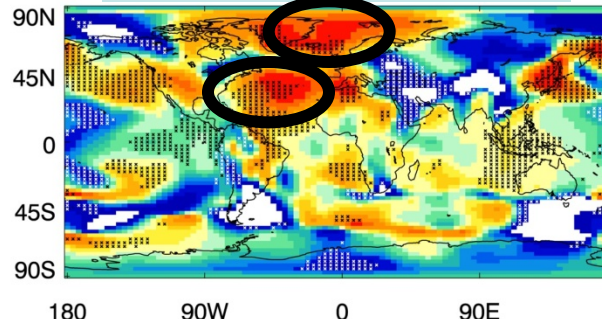
RPC in seasonal (DJF, top) & decadal (Yrs 2-5, bottom) hindcasts

Surface temperature



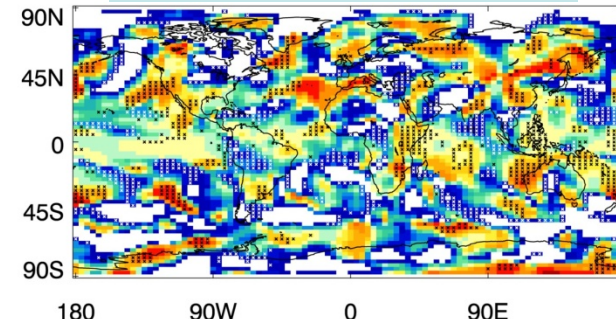
(d) SAT

Sea level pressure

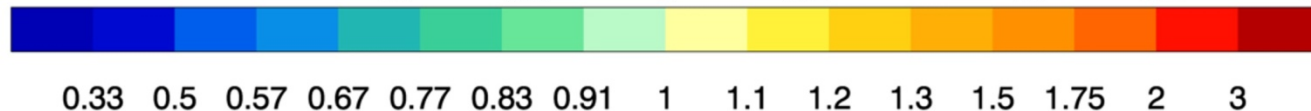
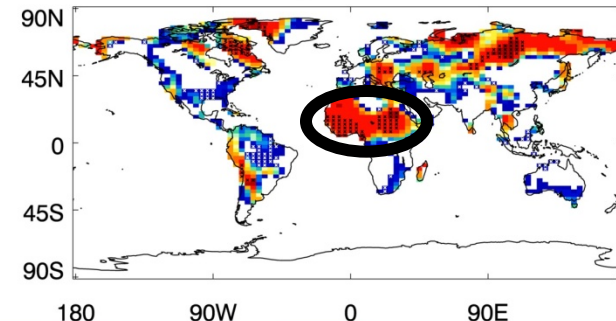
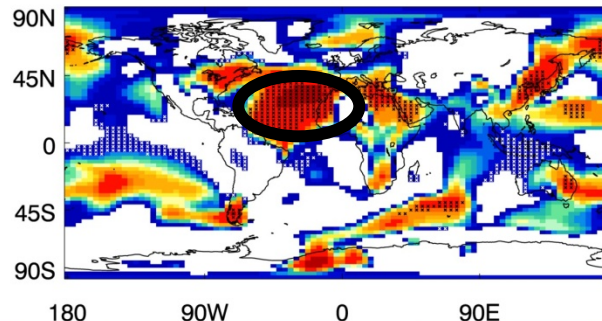
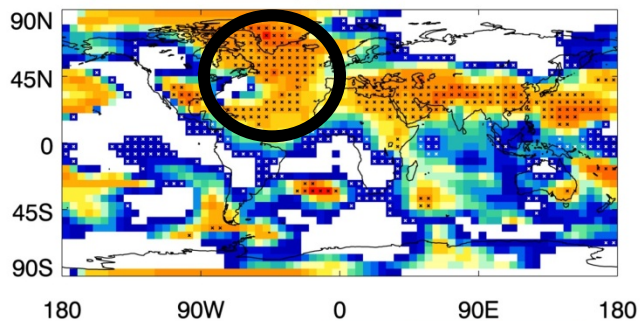


(e) MSLP

Precipitation



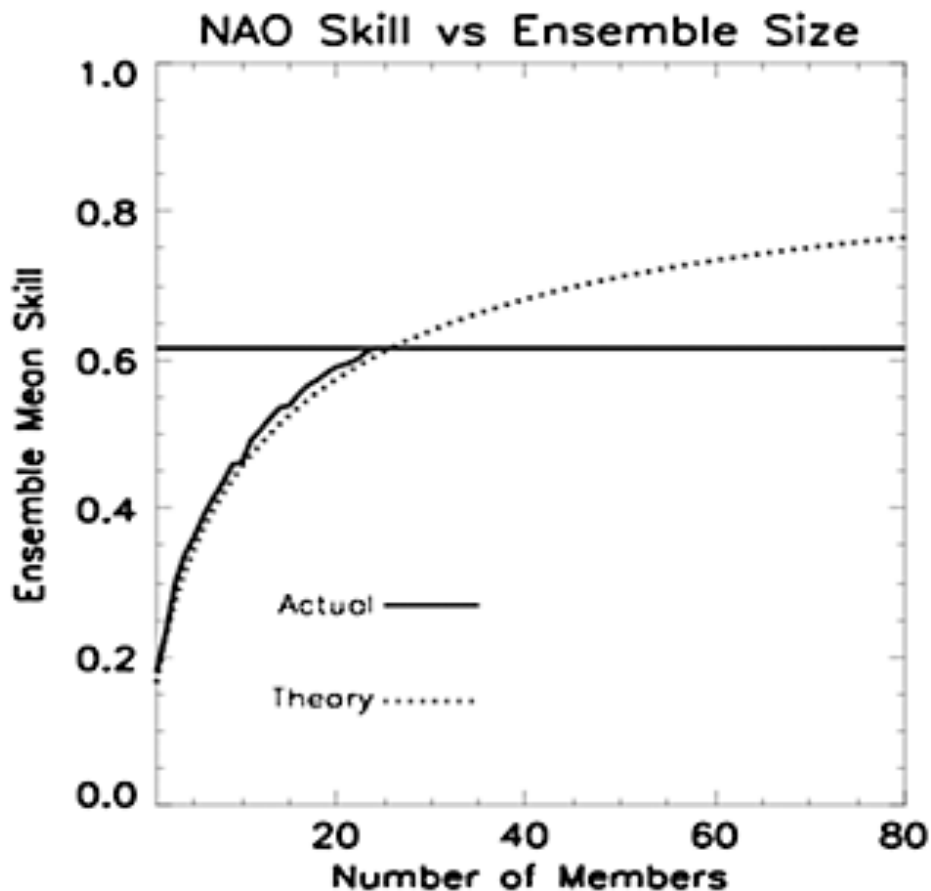
(f) PREC



RPC >> 1: Known regions of physical variability

NAO; North Atlantic decadal variability: MDR, SPG, Sahel

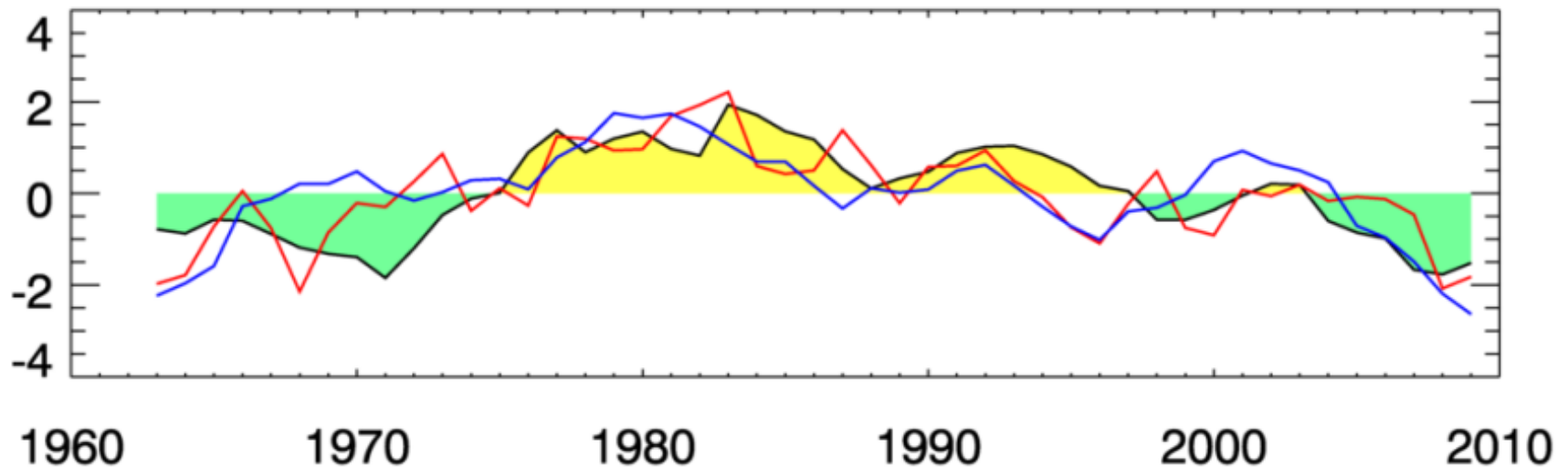
Effect of ensemble size on North Atlantic Oscillation skill



Scaife et al., 2014, GRL.

- Can now make skilful forecasts of NAO
- Correlation = 0.6
- But need large enough ensemble to get rid of noise
- Potentially even higher correlations with more members
- $RPC = 2.3 \gg 1$, i.e. under-confident (correlation not matched to signal-to-noise ratio)
- Need to correct variances to make $RPC = 1$

Can we predict the PDO?

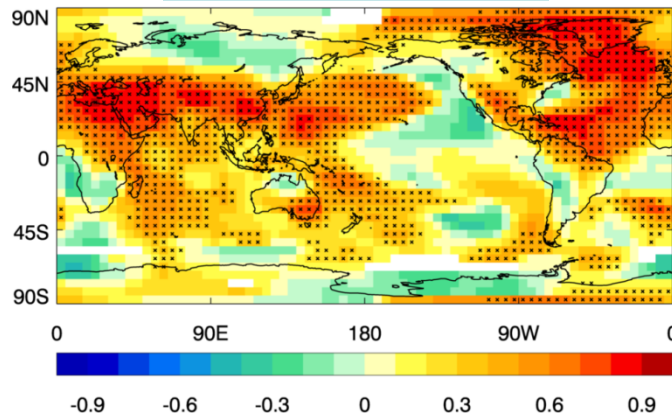


- Detrended 5-year running mean PDO
- CMIP5 decadal hindcasts (red) $r = 0.7$
 - yearly start dates (HadCM3, CanCM4, MPI, GFDL, MIROC5)
 - include 2 lagged start dates → 174 ensemble members
- equivalent uninitialized simulations (blue) $r = 0.6$
- **PDO potentially predictable but externally forced?**

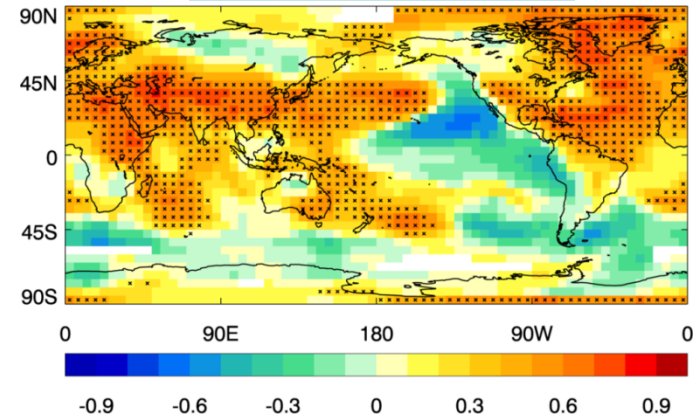
Can we predict the PDO?

Correlation of
detrended T
1960-2010

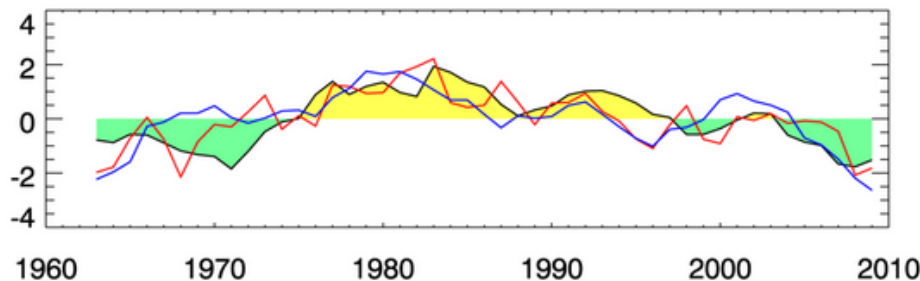
Initialized



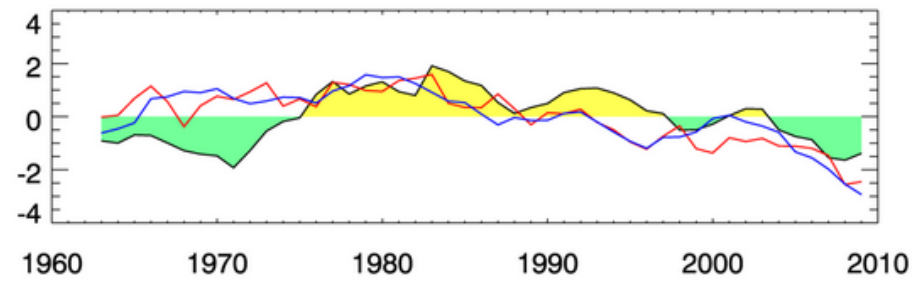
Uninitialized



Detrended PDO, $r=0.7$



Detrended PDO, $r=0.4$



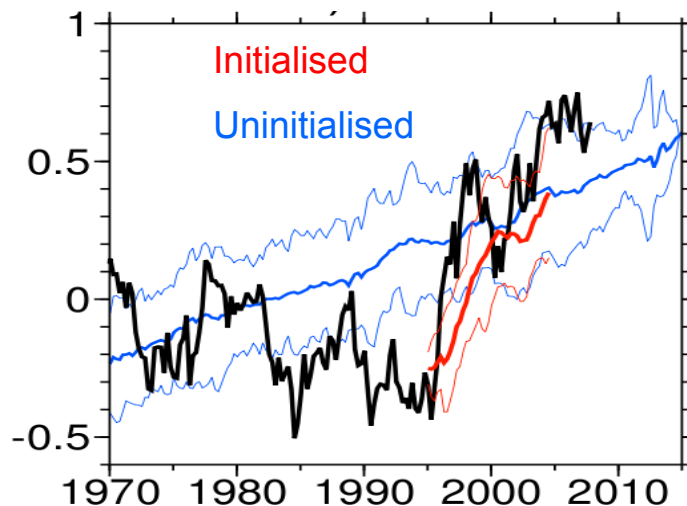
- No skill in eastern tropical Pacific (key region for the hiatus)
- Over longer period from 1880 **uninitialized skill drops to $r = 0.3$**
- **Apparent skill since 1960 not robust? And not eastern tropical Pacific**
- **Care needed when interpreting detrended skill!**



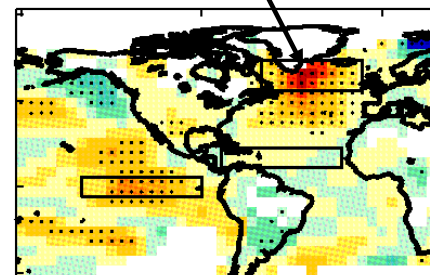
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Predictions of
sub-polar 1990s
gyre warming

Decadal skill: sub-polar gyre



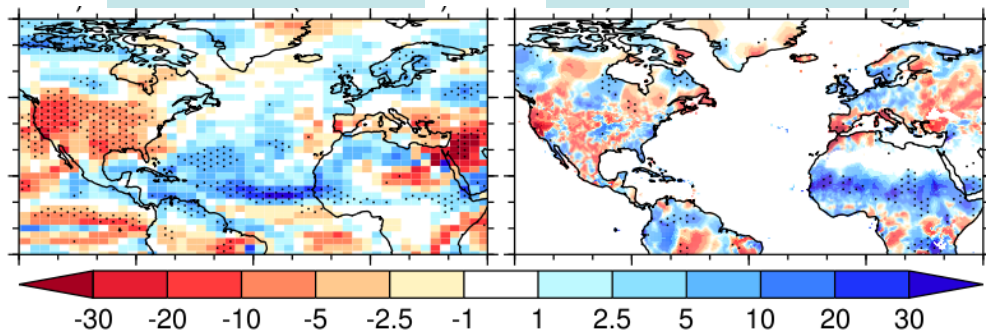
Sub-polar
gyre



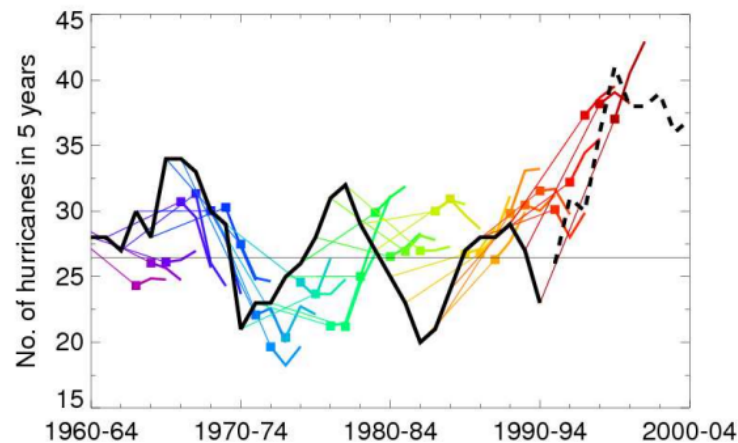
Impacts: rainfall

Model

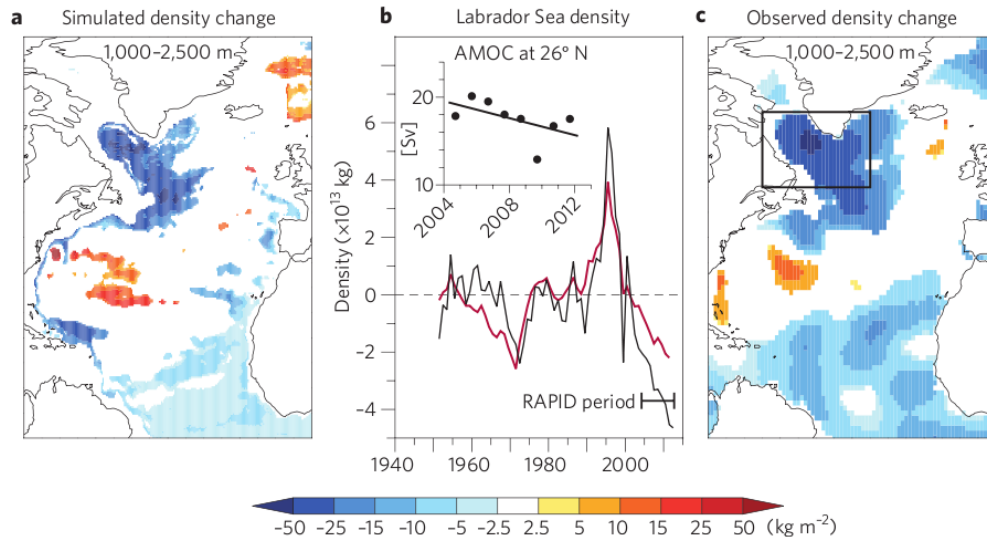
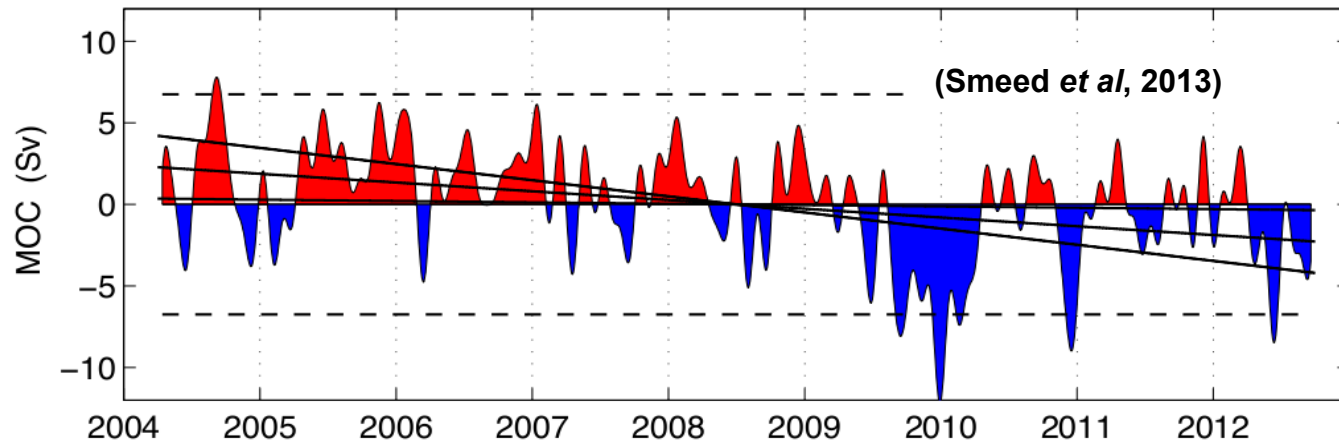
Observations



Impacts: hurricanes



Observed weakening of Atlantic overturning circulation



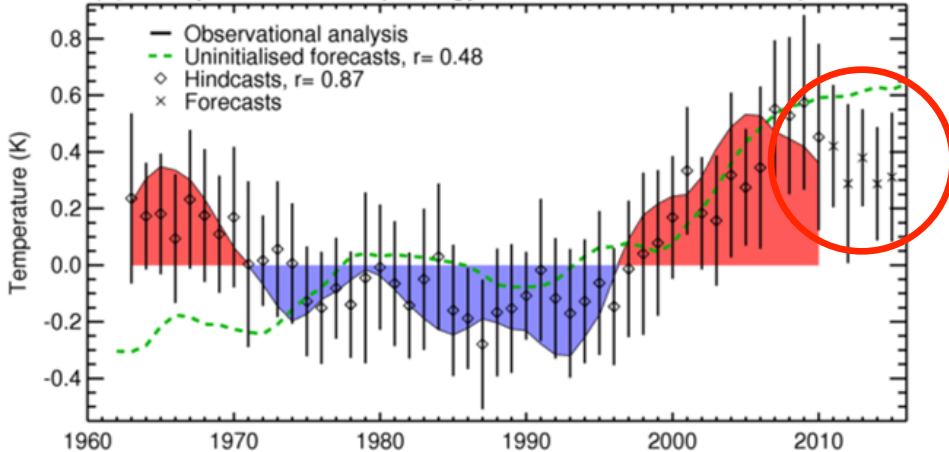
(Robson *et al*, 2013)

- Weakening of AMOC in direct observations (RAPID)
- and inferred from deep densities
- **0.5 Sv per year**
- **30% decrease in 10 years!**

Predicted cooling of North Atlantic

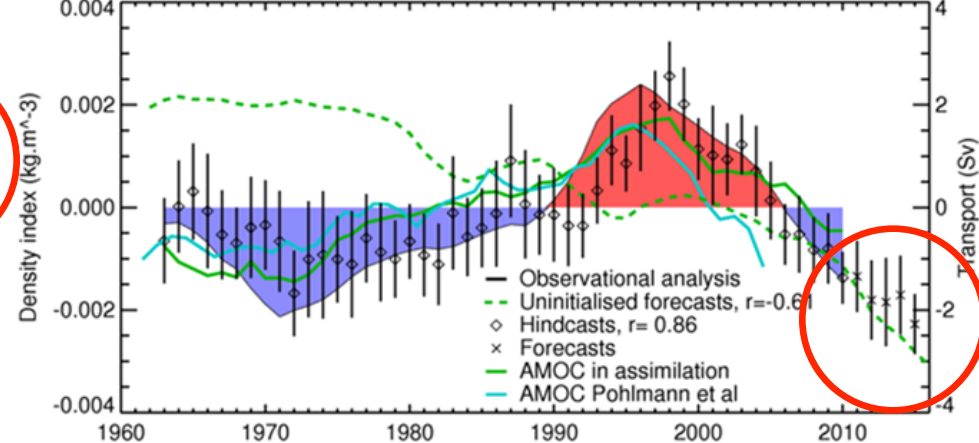
Temperature

(a) Five-year-mean subpolar gyre 500m mean ocean temperature



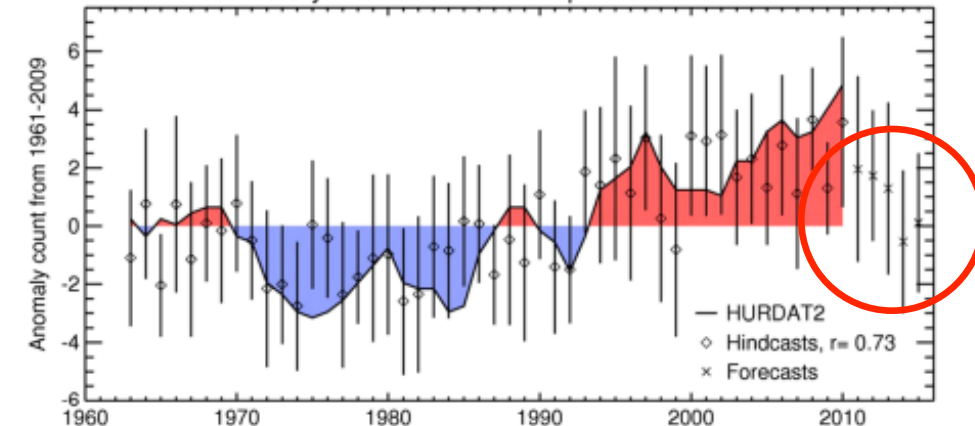
Ocean circulation

(c) Five-year-mean deep ocean density weighted difference index



Atlantic tropical storms

Five-year-mean Atlantic tropical storm count



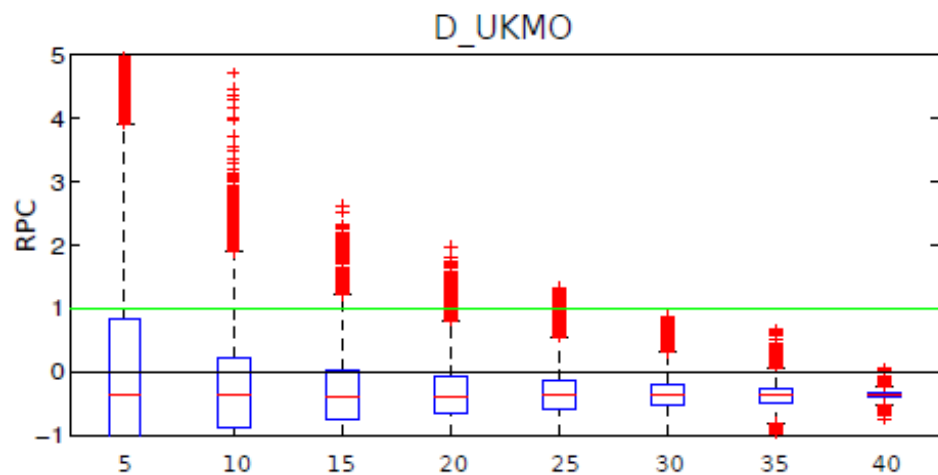
- Atlantic predicted to cool in response to weakening of ocean overturning
- Likely to cause climate impacts around the Atlantic basin
- Not a reversal, but impacts associated with warm Atlantic less likely:
 - cold winters and wet summers in Europe less likely
 - fewer hurricanes than recent peaks
 - reduced Sahel rainfall
 - reduced risk of drought in SW USA

Summary

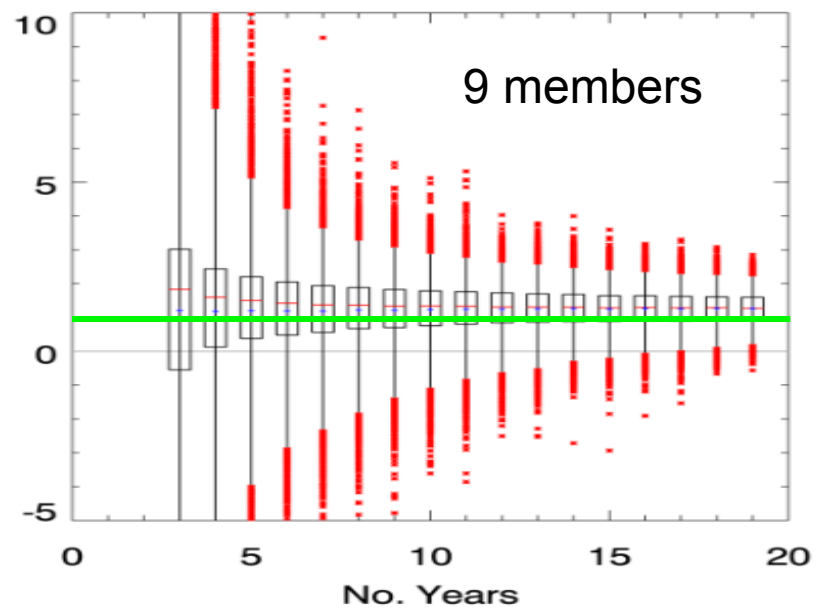
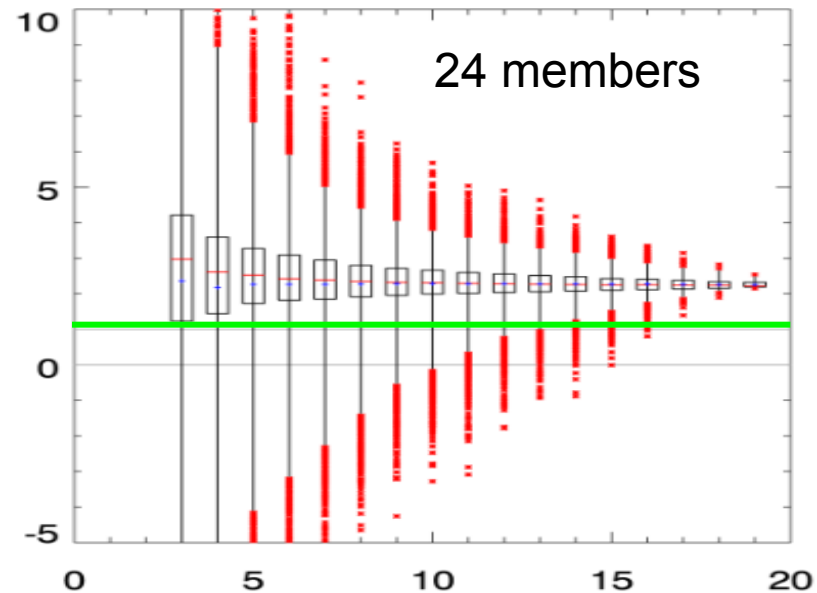
- Observed reduction in incoming energy at onset of hiatus
 - what is the role of ocean heat uptake?
- Response to Arctic sea ice is uncertain
- Models underestimate the predictability of the real world
 - need large ensemble and to adjust variance
 - some measures may underestimate skill
 - perfect model experiments not an upper limit!
- No skill for eastern tropical Pacific
- Potentially important changes in the Atlantic occurring now
 - Need to understand climate impacts
 - Multi-model forecasts → **DCPP component B !**

RPC sensitivity

GloSea3



GloSea5



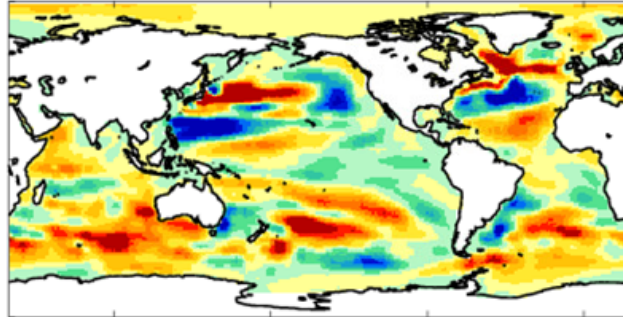
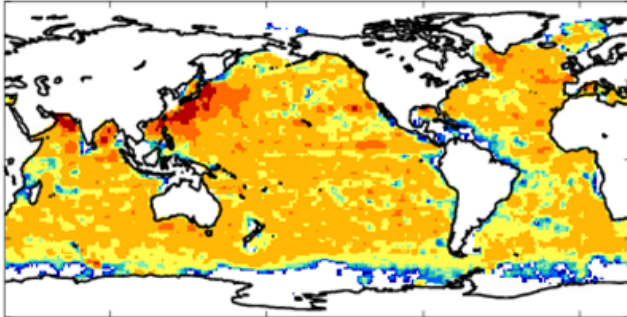
Assessment of Met Office ocean analysis

Number of observations

Heat content trend (H_t , Wm^{-2})

(a) All

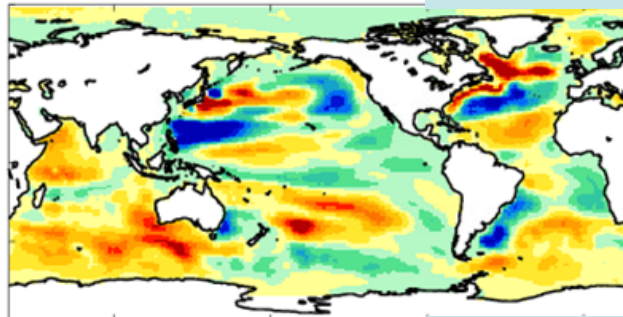
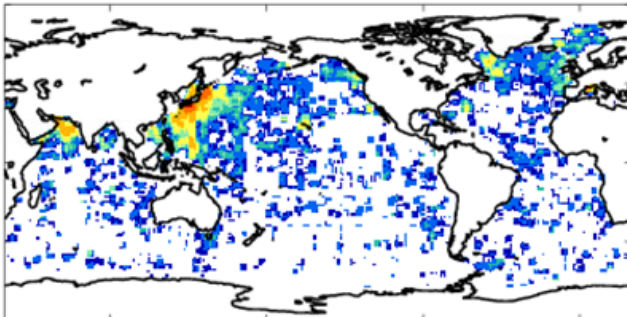
(b) All



(c) 1990s

(d) 1990s

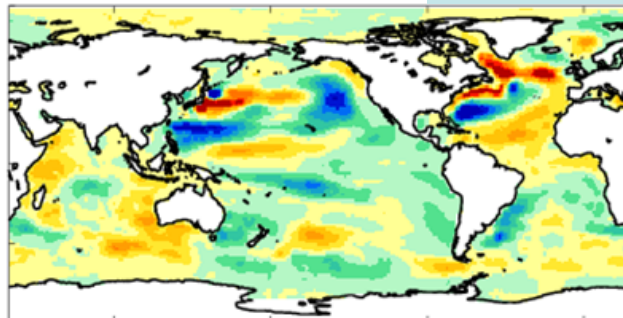
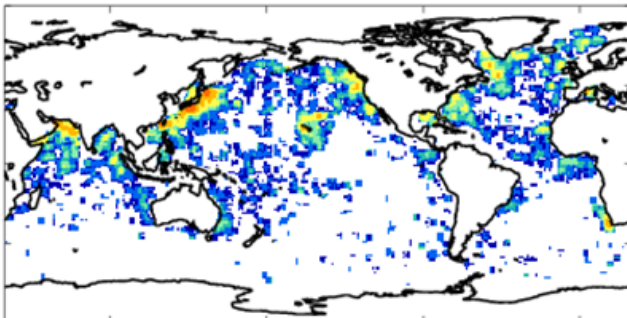
R=0.88



(e) 1960s

(f) 1960s

R=0.76



Data withholding experiments

- Reconstruct Argo period 2008-12 using sub-sampled observations typical of historical periods

- Uncertainties in annual mean global ocean heat content range from 28 ZJ using 1950-56 obs to 8 ZJ using 2005-11 obs

- Uncertainties in global H_t range from 0.7 Wm^{-2} using 1950-56 obs to 0.2 Wm^{-2} using 2005-11 obs

- Possible instrument errors and regions not adequately sampled are not included



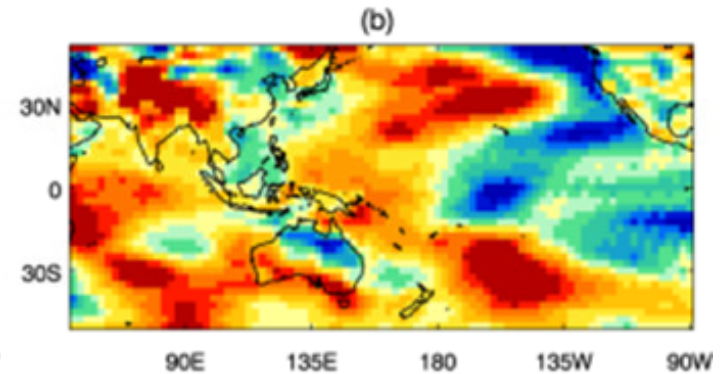
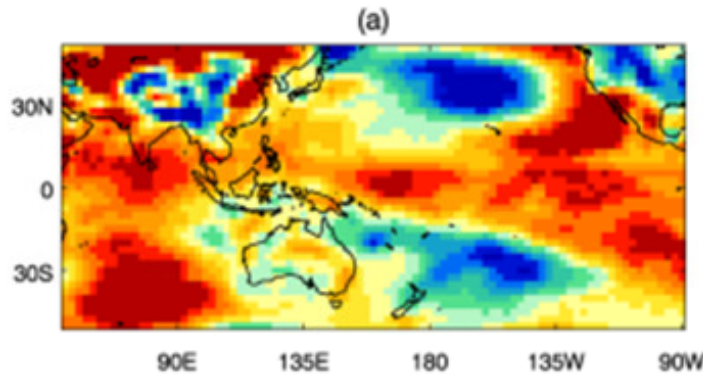
Met Office
Hadley Centre

Can we predict the PDO?

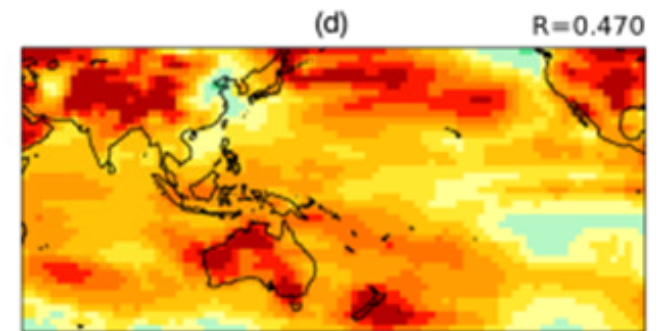
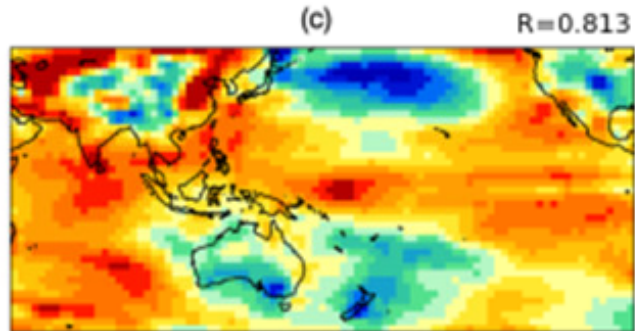
1970s shift

2000s shift

Observations



Initialized



Climatology

